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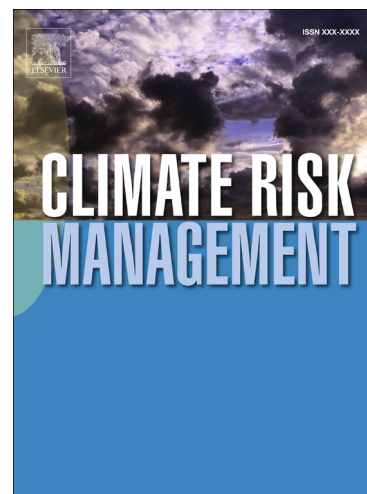
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1 **21st Century Flood Risk Projections at Select Sites for the U.S. National Park Service**

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8 9 **Abstract**

10 *Flood risk studies using stationary flood frequency analysis techniques is commonplace. However, it is*
11 *increasingly evident that the stationarity assumption of these analyses does not hold as anthropogenic*
12 *climate change could shift a site's hydroclimate beyond the range of historical behaviors. We employ*
13 *nonstationary flood frequency models using the generalized extreme value (GEV) distribution to model*
14 *changing flood risk for select seasons at twelve National Parks across the U.S. In this GEV model, the*
15 *location and/or scale parameters of the distribution are allowed to change as a function of time-variable*
16 *covariates. We use historical precipitation and modeled flows from the Variable Infiltration Capacity*
17 *model (VIC), a land-surface model that simulates land-atmosphere fluxes using water and energy*
18 *balance equations, as covariates to fit a best nonstationary GEV model to each site. We apply climate*
19 *model projections of precipitation and VIC flows to these models to obtain future flood frequency*
20 *estimates. Our model results project a decrease in flood risk for sites in the southwestern U.S. region and*
21 *an increase in flood risk for sites in northern and eastern regions of the U.S. for the selected seasons. The*
22 *methods and results presented will enable the NPS to develop strategies to ensure public safety and*
23 *efficient infrastructure management and planning in a nonstationary climate.*

24 **Declarations of Interest:** *none*

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37 1. Introduction

38 Anthropogenic climate change has increased global mean annual land-surface air temperatures and
39 evidence supports a change in the behavior of precipitation (Hartmann et al. 2013) and streamflow
40 extremes (Hirsch and Ryberg 2012; Mallakpour and Villarini 2015; Ahn and Palmer 2016). Given the non-
41 stationary nature of our climate system at present, the common assumption in traditional flood
42 frequency analysis techniques that flood risk will remain stationary into the future must be questioned -
43 climate change is anticipated to continue to shift hydroclimate beyond the range of historical behaviors
44 (Milly et al. 2008).

45 As temperatures rise, we expect an increase in total precipitable water in the atmosphere (Trenberth et
46 al. 2003), which was already observed over much of North America (Ross and Elliott 1996).
47 Consequently, Hartmann et al. (2013) suggest a likely observed increase in either the frequency or
48 intensity of heavy precipitation events across North America, particularly in central North America.
49 Studies using extreme value theory and precipitation-temperature scaling also generally support this
50 claim (DeGaetano 2009; Wasko and Sharma 2017).

51 However, trends in observed extreme streamflow are more variable (Ahn and Palmer 2016). Lins and
52 Slack (1999) found both increasing and decreasing trends in historical streamflow extremes in the
53 eastern U.S. with a general decrease in extremes in western U.S., the Pacific Northwest, and the
54 Southern Plains. Mallakpour and Villarini (2015) found an increase in the frequency of observed floods
55 in the central U.S., with no evidence to support a change in the observed magnitude of flood events. In
56 the southwestern U.S., Hirsch and Ryberg (2012) found decreasing flood magnitudes associated with
57 increasing atmospheric greenhouse gas (GHG) levels, while the eastern and northeastern U.S. showed
58 increasing, but non-significant, flood magnitude trends in response to carbon dioxide increases.

59 Flood risk analysis using distributions like the log-Pearson type III (LPIII) distribution, generalized
60 extreme value (GEV) distribution, generalized Pareto distribution (GPD), and lognormal distribution, all
61 of which assume stationarity of risk, is commonplace (Stedinger et al. 1993; Coles 2001; England et al.
62 2018). Several more recent approaches assess time-varying (i.e., nonstationary) characteristics of flood
63 risk. AghaKouchak et al. (2013) and Salas et al. (2018) provide a detailed review of nonstationary
64 extreme value analysis methods. Applying a nonstationary GEV distribution and allowing the location
65 and/or scale of the distribution to change linearly as a function of time or various hydrometeorological
66 covariates is one approach to assess changing flood risk (Coles 2001; Salas and Obeysekera 2014;
67 Condon et al. 2015). This framework has been applied to extreme streamflow using time (Katz et al.
68 2002; Salas and Obeysekera 2014), meteorological variables (Towler et al. 2010; Condon et al. 2015),
69 and climate indices (Lima et al. 2015) as covariates. Further, Condon et al. 2015 assessed future flood
70 risk with this model framework using future projections of covariates generated from global climate
71 models (GCMs).

72 For 12 National Park Service (NPS) sites (chosen to capture an array of hydroclimates in the U.S.) we
73 project future 21st century flood risk by applying the nonstationary generalized extreme value
74 distribution and projections of hydrometeorological variables from an ensemble of GCMs covering two
75 Representative Concentration Pathways (RCP). There are few applications of nonstationary flood risk
76 analysis to the management of U.S. public lands and conservation areas - the results presented in this
77 work will help enable the NPS to better understand flood risks in a nonstationary context, which could
78 subsequently be used for efficient short- and long-term management of protected resources.

79 **2. Methods**80 ***Nonstationary Generalized Extreme Value Distribution***

81 The starting point of the nonstationary flood frequency model is the assumption that the seasonal or
 82 annual flow extremes are assumed to follow the generalized extreme value distribution, a common
 83 statistical tool used in hydrological extreme value analysis. Described with further detail in Coles (2001),
 84 block maxima of independent and identically distributed random variables follow the generalized
 85 extreme value distribution, with the cumulative distribution function:

$$G(z) = \exp \left\{ - \left[1 + \varepsilon \left(\frac{z - \mu}{\sigma} \right) \right]^{\frac{-1}{\varepsilon}} \right\} \quad (1)$$

86 where $\{z: 1 + \varepsilon(z - \mu)/\sigma \geq 0\}$. The variable z is the streamflow maxima and the parameters μ , σ , and
 87 ε represent the distribution location, scale, and shape, respectively. The location determines the
 88 position of the distribution, the scale determines the spread of the distribution, and the shape
 89 determines the behavior of the upper tail. Equation (1) follows the form of the type I extreme value
 90 distribution (EVI), or Gumbel distribution, when the shape (ε) is 0 (light tail). Similarly, equation (1)
 91 follows the form of the EVII, or Frechet distribution, when the shape (ε) is positive (heavy tail) and the
 92 EVIII, or Weibull distribution, when the shape is negative (bounded tail). Coles (2001) provides details on
 93 extreme value theory.
 94

95 Nonstationarity is incorporated by allowing the location or both the location and scale parameters of
 96 equation (1) to vary as a function of covariates. The nonstationary location and scale are modeled as
 97 follows:

$$98 \quad \mu(t) = \beta_{0,\mu} + \beta_{1,\mu}x_{1,t} + \dots + \beta_{n,\mu}x_{n,t} \quad (2)$$

$$99 \quad \sigma(t) = \exp(\beta_{0,\sigma} + \beta_{1,\sigma}x_{1,t} + \dots + \beta_{n,\sigma}x_{n,t}) \quad (3)$$

100 where x variables represent covariates and β denotes the fitted parameters. The transformed scale
 101 parameter is used to ensure the scale is positive. Stationary and nonstationary GEV parameters are
 102 estimated using the method of maximum likelihood (MLE), a general and flexible parameter estimation
 103 technique also used in similar studies (Katz et al. 2002; Towler et al. 2010; Condon et al. 2015).
 104

105 The best nonstationary model (i.e. the best set of covariates) is selected by minimizing the Akaike
 106 Information Criteria (AIC), which penalizes the negative maximized log-likelihood of a model for the
 107 number of parameters used. AIC is defined by:

$$108 \quad AIC = 2(NLLH) + 2(k) \quad (4)$$

109 where $NLLH$ is the negative maximized log-likelihood obtained from MLE and k is the number of
 110 independently adjusted model parameters (Akaike 1998). As an alternative to AIC, similar nonstationary
 111 GEV studies have used the likelihood ratio test, a common statistical tool used to test the significance of
 112 improvement in maximized log-likelihoods for nested models. However, with the number of models we
 113 test for in this study, outcomes of the likelihood ratio test would lose their interpretability (Katz 2013)

114 and some of the models we fit are not nested. For this reason, the likelihood ratio test is not used as the
 115 primary selection criteria, though nonstationary models selected by AIC are still compared with the
 116 stationary GEV distribution with the likelihood ratio test.

117 Exceedance probability levels for stationary GEV distributions are solved with equation (5):

$$118 \quad z_p = \mu - \frac{\sigma}{\varepsilon} \left[1 - \left\{ -\ln(1-p) \right\}^{-\varepsilon} \right] \quad (5)$$

119 where z_p is the streamflow with exceedance probability p and the parameters μ , σ , and ε represent the
 120 GEV distribution location, scale, and shape (with $\varepsilon \neq 0$) (Coles 2001). Traditional stationary return level
 121 calculations are not applicable in a nonstationary context, where exceedance periods change with each
 122 new GEV distribution. We follow the methods explained in Salas and Obeysekera (2014) and Condon et
 123 al. (2015) for nonstationary risk assessment. The above methods were largely implemented in R (R Core
 124 Team 2016) with the package 'extRemes' (Gilleland and Katz 2016).

125 For comparison to the stationary and nonstationary GEV models, we also fit a stationary log-Pearson
 126 type III distribution to flow maxima. LPIII distributions are fit using the method of moments following
 127 USGS Bulletin #17B flood flow frequency guidelines (IACWD 1982). We include a visual process summary
 128 of stationary and nonstationary GEV flood frequency analysis in pages i-iii in Appendix A.

129 Here we assess future flood risk using an ensemble of climate model outputs (further described in
 130 subsequent sections). We first select a best nonstationary GEV distribution from a set of observed
 131 covariates. We then simulate model behavior with an ensemble of climate models to evaluate the risk of
 132 exceeding some site-specific critical flow within a selected design life. Steps for the analysis are:

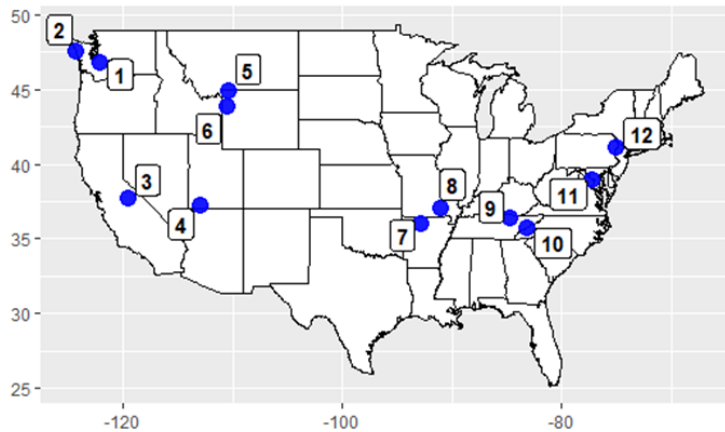
- 133 1. A performance period of interest (e.g., 2040-2069), a project life (e.g., 20 years), and a critical flow are
 134 selected for a site.
- 135 2. One climate model is selected at random from the ensemble of climate models. From the randomly
 136 selected model, a block of covariate data is randomly selected within the period of interest and with a
 137 length of the project life (e.g., a 20-year block of data is selected from 2040-2069 model data).
- 138 3. The best nonstationary GEV distribution is applied to the selected block of covariate data to
 139 determine year-specific risks of exceeding the critical flow.
- 140 4. Following Salas and Obeysekera (2014), the total risk of exceeding the critical flow within the project
 141 life is calculated (e.g., the risk of exceeding the critical flow over the 20-year project life).
- 142 5. Steps 2-4 represent one simulation. This process is repeated for each RCP scenario, multiple climate
 143 models and the many blocks of covariate data with a length of the project life within the period of
 144 interest. This provides a distribution of simulated probabilities of exceeding the critical flow over the
 145 project life.

146 **3. Study Sites and Data**

147 ***Study Sites***

148 Twelve USGS streamflow gauge sites of interest to the NPS are the focus of this study. Figure 1 provides
 149 details regarding the sites and their locations. These sites have a long historical USGS gauging record and

150 represent a diverse array of hydroclimates where impactful flooding events occurred in the past.
 151 Further, these basins contain minimal hydrologic alteration, ensuring that human-caused land cover
 152 change and river alterations (e.g., diversions, dams, and other structures) are not impacting these study
 153 sites. Some recent notable and documented flood events for these basins include the January 1997
 154 flood in Yosemite National Park, 2006 flooding in Mount Rainier National Park, and the 2017 flooding in
 155 the Ozark National Scenic Riverways.

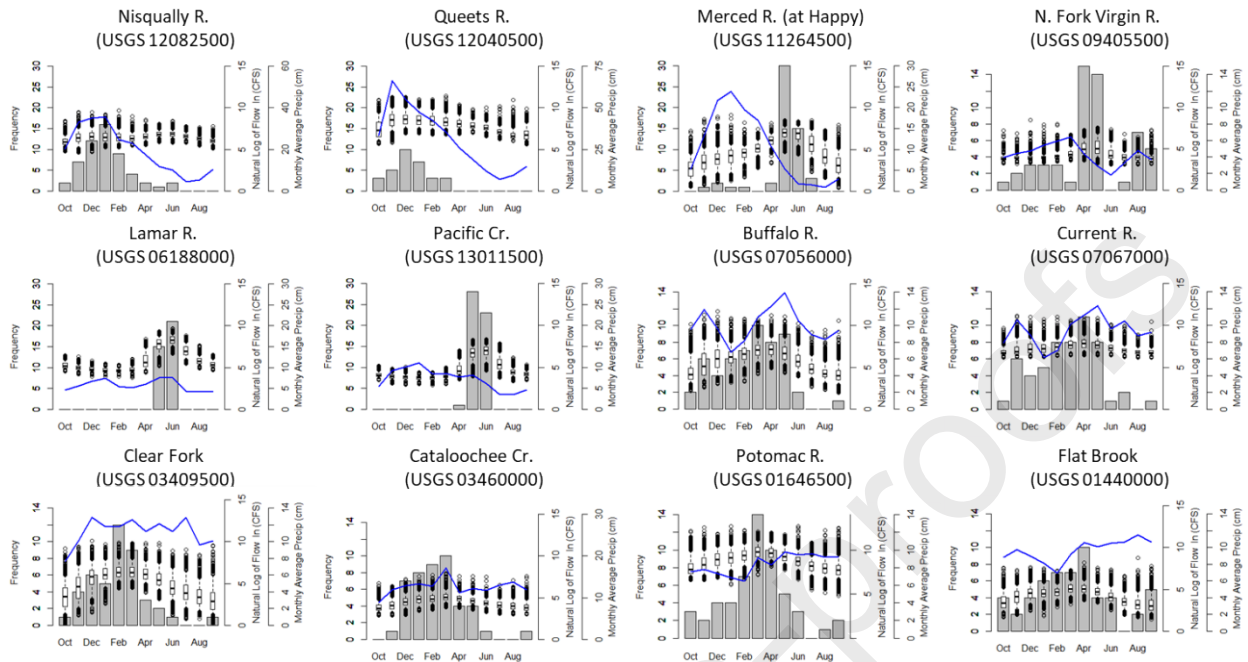


	USGS Site	USGS Site Description	Unit	Drainage Area (sq mi)
1	12082500	Nisqually River near National, WA	Mount Rainier National Park	133
2	12040500	Queets River near Clearwater, WA	Olympic National Park	445
3	11264500	Merced River at Happy Isles Bridge near Yosemite, CA	Yosemite National Park	181
4	09405500	North Fork Virgin River near Springdale, UT	Zion National Park	344
5	06188000	Lamar River near Tower Ranger Station, YNP	Yellowstone National Park	668
6	13011500	Pacific Creek at Moran, WY	Grand Teton National Park	169
7	07056000	Buffalo River near St. Joe, AR	Buffalo National River	829
8	07067000	Current River at Van Buren, MO	Ozark National Scenic Riverways	1,667
9	03409500	Clear Fork near Robbins, TN	Big South Fork National River and Recreation Area	272
10	03460000	Cataloochee Creek near Cataloochee, NC	Great Smoky Mountains National Park	49
11	01646500	Potomac River near Washington, D.C. Little Falls Pump	Chesapeake & Ohio Canal National Historical Park	11,560
12	01440000	Flat Brook near Flatbrookville, NJ	Delaware Water Gap National Recreation Area	64

156
 157 *Figure 1 Location (top) and descriptions (below) of the 12 sites.*

158 Drainage areas of the selected basins range from 49 to 11,560 square miles. The sites have varied
 159 characteristics in terms of the timing of annual maxima, monthly precipitation, and streamflow
 160 seasonality, as shown in Figure 2. Sites in the northwest (Nisqually R. and Queets R.) experience flood
 161 events during the winter wet season. Western sites (Merced R., North Fork Virgin R., Lamar R., and
 162 Pacific Cr.) exhibit delayed spring streamflow response to winter precipitation, suggesting snowmelt
 163 driven river systems. Similarly, historical flooding events often occur in the spring for these sites,
 164 suggesting snowmelt might be an important driving mechanism for flooding events at these sites. The
 165 remaining eastern U.S. sites (Buffalo R., Current R., Clear Fork, Cataloochee Cr., Potomac R., and Flat
 166 Brook), exhibit variable streamflow and precipitation characteristics, with the majority of floods
 167 clustered over October-June.

168



169

170 *Figure 2 For each study site a bar plot for the count of annual peak mean daily flows occurring within each month (left axis),*
 171 *boxplots of mean daily flows for each month (right axis), and a line plot of average monthly precipitation (far right axis) using*
 172 *1951-2005 data.*

173 The months we use for seasonal analysis at each site are based on the timing of annual peak flows
 174 (water year), the distribution of daily flows for each month, and the monthly average precipitation.
 175 Generally, the season we select for analysis includes consecutive months that experience the highest
 176 frequency of annual maximum mean daily flows. We also assess monthly average precipitation and daily
 177 streamflow patterns to assess potential dominant flood mechanisms (e.g., runoff and snowmelt flood
 178 drivers), and we consider historical trends in the timing of observed seasonal peak flows. As further
 179 described in the coming section, the season we select to investigate for each site also corresponds to
 180 the seasonal covariates we use. To capture antecedent conditions that might influence flooding (e.g.,
 181 snowpack), we also include covariates from the previous season.

182 **Data**

183 We use observed USGS gauge mean daily streamflow measurements available between 1951 and 2005
 184 (water year) for analysis (U.S. Geological Survey 2016). Water years missing data within the season of
 185 interest are excluded from the analysis.

186 We use 1951-2005 (water year) observed season average daily precipitation of each contributing basin
 187 and season average daily hydrologic model generated flow as covariates – daily values of both are
 188 provided by the U.S. Bureau of Reclamation. These are determined using Livneh et al. (2015) 1/16°
 189 spatially gridded meteorological data derived from NOAA Cooperative Observer Network stations.
 190 Hydrologic model flows provided by the U.S. Bureau of Reclamation are generated from the Variable
 191 Infiltration Capacity model (VIC). VIC is a land-surface model that simulates spatially gridded, land-
 192 atmosphere fluxes using the water and energy balance equations (Liang et al. 1994). Modeled flows are

193 generated using VIC version 4.1.2h. This model requires daily precipitation, maximum and minimum air
194 temperature, and wind speed as input forcings required (Prata 1996; Kimball et al. 1997; Thornton and
195 Running 1999; Bohn et al. 2013). We use land-cover input data and calibrated parameters from Maurer
196 et al. (2002) and Livneh et al. (2013). Details on the VIC model are available in:
197 <http://vic.readthedocs.io/en/master/>. VIC river routing was performed at 1/16° grids using the routing
198 model from Lohmann et al. (1996).

199 The use of VIC model flows as a covariate in nonstationary flood frequency analysis is a novel
200 contribution of this research. We introduce this because we posit that VIC model flows better capture
201 the water and energy balance features of a basin as well as basin specific land-cover features compared
202 to average meteorological covariates (e.g., precipitation). We assess and summarize VIC model
203 performance compared to observed flows for each site and season in Table A-1; while the modeled
204 flows for several of the sites have strong biases, the correlations between VIC model flows and observed
205 streamflow for each site are strong. While the observed magnitude of daily flow (and potentially the
206 observed magnitude of the seasonal peak daily flow) might be poorly captured by VIC model
207 simulations, the seasonal average flow from the VIC model corresponded well with the observations and
208 thus, is a valuable covariate. Furthermore, we found a strong correlation between seasonal average
209 flows from the VIC model and the peak mean daily flow for the season of interest for each site (Table A-
210 2). This suggests the seasonal average flows contain information about the seasonal peak flow; also, the
211 VIC model flows capture the hydrologic processes in the basin providing complementary information.
212 With this motivation, we use the seasonal average flows from VIC model as one of the covariates in the
213 nonstationary GEV model.

214 We use an ensemble of projected 1951-2099 (water year) season average daily precipitation of each
215 contributing basin and season average daily VIC model generated flow as future covariates, which
216 enables 21st century projections of flood risk. Daily values of both are provided by the U.S. Bureau of
217 Reclamation. Projections are determined using the U.S. Bureau of Reclamation's LOCA CMIP5 dataset.
218 This dataset contains 64 projections of daily, 1/16° gridded precipitation and maximum/minimum
219 temperature from an ensemble of 32 general circulation models, covering two different greenhouse gas
220 RCPs. We investigate RCP 8.5, a scenario representing high and increasing greenhouse gas levels into the
221 future, and RCP 4.5, a scenario representing a radiative forcing stabilization scenario (van Vuuren et al.
222 2011). LOCA CMIP5 data is generated from bias corrected and downscaled coarse GCM data (with a
223 spatial resolution generally exceeding 1°) from the CMIP5 multi-model ensemble (Taylor et al. 2011).
224 Additional information on these processes can be found in Pierce et al. (2014, 2015) and Reclamation
225 (2016). This data is available from the downscaled CMIP3 and CMIP5 climate and hydrology projections
226 archive at https://gdo-dcp.ucllnl.org/downscaled_cmip_projections/. Information on the CMIP5 project
227 can be found in Taylor et al. (2011). The GCMs we use through the LOCA CMIP5 dataset, the responsible
228 modeling groups, and an acknowledgement of the World Climate Research Program's Working Group on
229 Coupled Modelling are presented in Table A-3. The same methods as described earlier are used to
230 generate VIC model flows. However, because average daily wind speed is not available in the LOCA
231 CMIP5 dataset, historical Livneh et al. (2015) daily average wind speeds are used for the projected VIC
232 wind speed forcing.

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241 **Overview**

242 A process summary of this research appears in Figure 3. For each site of interest, we first fit a best
243 nonstationary GEV model to observed historical seasonal peak flows considering historical season
244 average and previous season average daily precipitation and hydrologic model generated flows as
245 potential covariates. For each site a set of models is generated by fitting nonstationary GEV distributions
246 to different combinations of these covariates, and, as mentioned, the best model (i.e., the best subset of
247 covariates) is selected using AIC. We then apply to the best model for each site the LOCA CMIP5
248 ensemble of future covariate projections through 2099 (water year). For each year of each GCM
249 ensemble member, the GEV distribution is projected using the projected covariate values. This provides
250 time-varying estimates of flood frequency distributions into the future. We also include traditional flood
251 frequency models (stationary GEV and LPIII distributions) in our analysis for comparison.

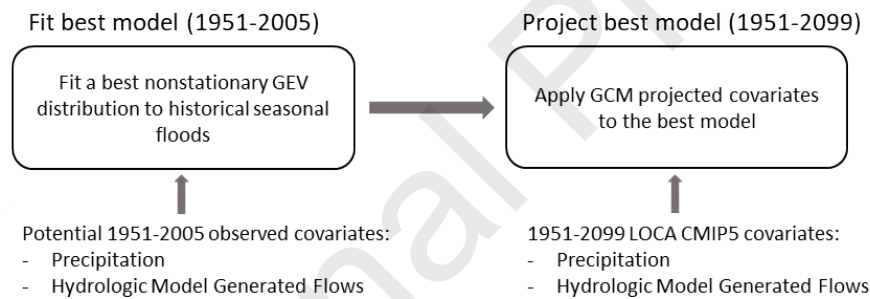


Figure 3 Process structure of this study.

252 **4. Results**

253 The best nonstationary GEV model was evaluated for each site using observed season and previous
254 season average precipitation and VIC model flow as potential covariates. Table 1 lists the covariates
255 selected in the best model for each park.

	USGS Site	USGS Site Description	Unit	Season of	Previous Season	GEV Location	GEV Scale	P-Value*
				Analysis	Covariates			
1	12082500	Nisqually River near National, WA	Mount Rainier National Park	Oct-Mar	Oct-Dec	VIC	VIC	3.8E-08
2	12040500	Queets River near Clearwater, WA	Olympic National Park	Oct-Mar	Oct-Dec	VIC		9.4E-05
3	11264500	Merced River at Happy Isles Bridge near Yosemite, CA	Yosemite National Park	Apr-Jun	Nov-Mar	Pre_Pr + Pr		<2.2E-16
4	09405500	North Fork Virgin River near Springdale, UT	Zion National Park	Apr-Jun	Nov-Mar	VIC	VIC	1.6E-13
5	06188000	Lamar River near Tower Ranger Station, YNP	Yellowstone National Park	Apr-Jun	Nov-Mar	Pre_VIC + VIC	Pre_VIC + VIC	8.8E-06
6	13011500	Pacific Creek at Moran, WY	Grand Teton National Park	Apr-Jun	Nov-Mar	Pre_Pr + Pr		4.2E-10
7	07056000	Buffalo River near St. Joe, AR	Buffalo National River	Jan-May	Oct-Dec	Pr	Pr	8.4E-10
8	07067000	Current River at Van Buren, MO	Ozark National Scenic Riverways	Feb-May	Nov-Jan	Pre_VIC + Pr	Pre_VIC + Pr	3.4E-11
9	03409500	Clear Fork near Robbins, TN	Big South Fork National River and Rec. Area	Dec-Mar	Oct-Nov	Pr		8.7E-07
10	03460000	Cataloochee Creek near Cataloochee, NC	Great Smoky Mountains National Park	Dec-Mar	Oct-Nov	VIC	VIC	2.0E-05
11	01646500	Potomac River near Washington, D.C. Little Falls Pump	Chesapeake & Ohio Canal National Hist. Park	Feb-May	Nov-Jan	VIC	VIC	8.7E-10
12	01440000	Flat Brook near Flatbrookville, NJ	Delaware Water Gap National Rec. Area	Feb-Apr	Nov-Jan	Pre_VIC + VIC	Pre_VIC + VIC	4.7E-09

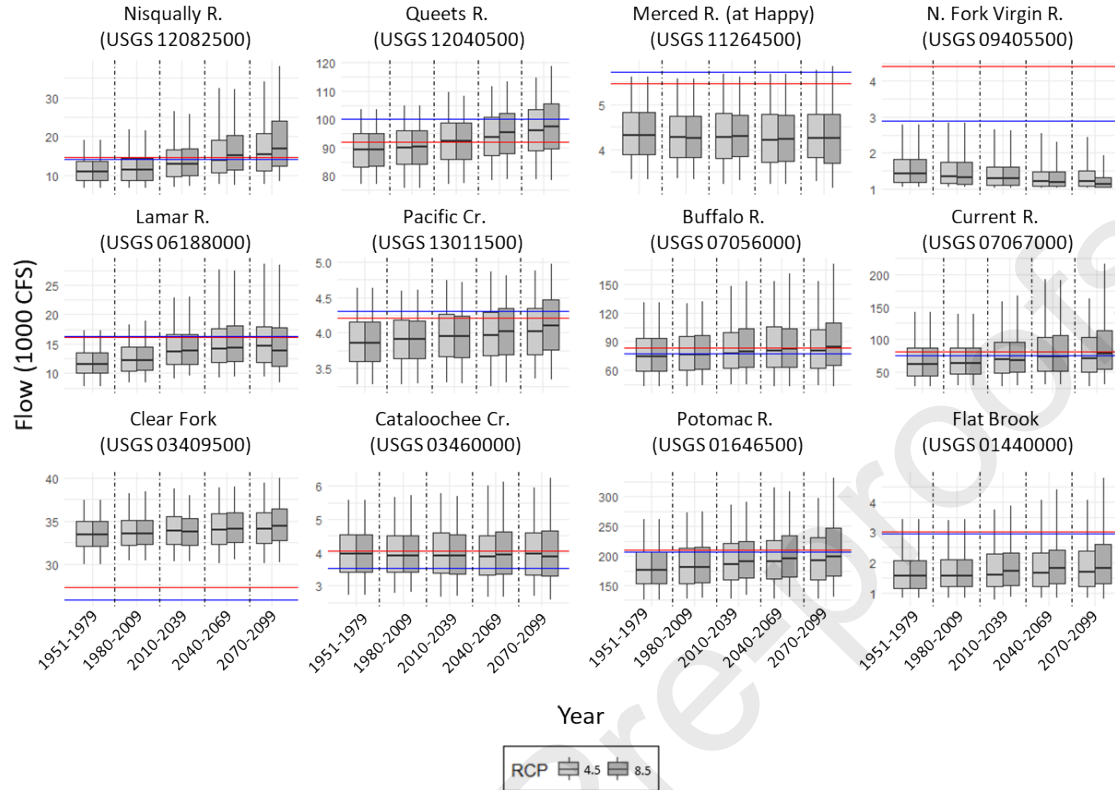
256

257 *Table 1 Best model parameters selected for the 12 sites. 'Pr' represents the seasonal average precipitation covariate and 'VIC'*
 258 *represents seasonal average flow covariate. A 'Pre' prefix indicates a previous season average covariate.*

259 For all sites, nonstationary GEV distributions are selected over stationary GEV distributions based on AIC
 260 scores. The p-values from the likelihood ratio test (compared to the stationary GEV distribution) are
 261 also included, all of which are less than 0.05. VIC flows are selected as a covariate in the best model for
 262 eight of the sites while precipitation is selected in the best model for five of the sites. Five models select
 263 covariates from the previous season. The best models for eight of the sites have both a nonstationary
 264 location and scale. The remaining best models only have a nonstationary location, these are the sites
 265 with a blank in the 'GEV Scale' column of Table 1. Appendix B includes more detail on the site specific
 266 fitted parameter values for each best model. For all but two sites, the location and/or scale of the
 267 models shift upward with an increase in the selected covariate; higher previous season VIC flows for
 268 USGS 06188000 (Lamar River) and 01440000 (Flat Brook) result in a decrease in the GEV location and/or
 269 scale parameters. One possible explanation of this is that both of these sites experience winter
 270 snowfall, so a higher historical previous season VIC flow for these sites might suggest earlier winter
 271 snowmelt which will decrease the likelihood of obtaining high spring peak flows.

272 For each site, we use the best nonstationary GEV model (Table 1) and the ensemble of LOCA CMIP5
 273 covariate projections to obtain an ensemble of 1% seasonal exceedance probability flows from 1951 to
 274 2099 (water year). Because there are 64 model runs in the LOCA CMIP5 ensemble, 64 1% exceedance
 275 probability flows are generated for every year (32 for RCP 4.5 and 32 for RCP 8.5). 1% exceedance
 276 probability flows for each RCP scenario are grouped into approximately 30-year time periods from 1951
 277 to 2099 (water year) and box plotted. The results for all sites appear in Figure 4. The 1% seasonal
 278 exceedance probability flows generated from the stationary LPIII distribution (blue line) and stationary
 279 GEV distribution (red line) fit to historical observed floods are also included. Similar plots for the 2% and
 280 0.2% exceedance probability flows for each site are available in Appendix B.

GCM Ensemble 1% Seasonal Exceedance Probability Flows



281

282 *Figure 4* Boxplots of 1% exceedance probability flows generated from the best GEV model and LOCA CMIP5 covariate projections
 283 for each site. Stationary LPIII (blue line) and GEV (red line) 1% exceedance probability levels are also included.

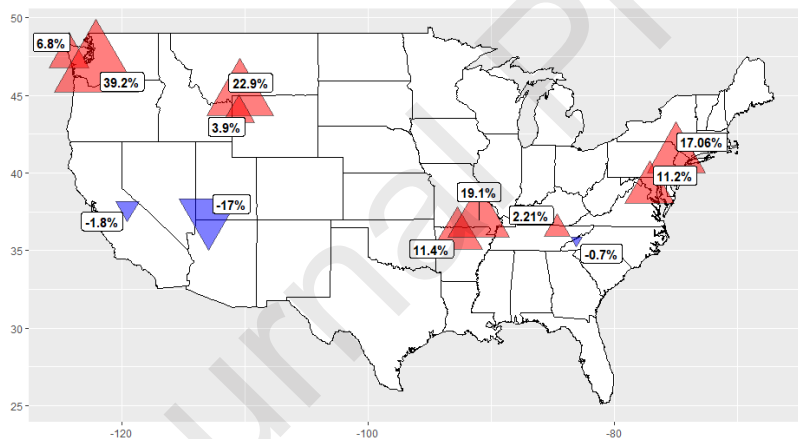
284 For some sites (Nisqually R. and Queets R., for example), an increase in all quantiles of 1% exceedance
 285 probability flows generated from the LOCA CMIP5 ensemble into the future is apparent (shown as an
 286 upward shift in the boxplots over time). The opposite is apparent for USGS 09405500 (N. Fork Virgin R.),
 287 which is showing decreasing trends. For many of the sites (Buffalo R. and Current R., for example), we
 288 see an increase in the interquartile range and an increase in the difference between the 5th and 95th
 289 percentiles of the ensemble 1% exceedance probability flows. For a site like USGS 03460000
 290 (Cataloochee River), where the median remains relatively steady, the changes in the interquartile range
 291 and 5th and 95th percentiles suggests an increase in variability in the magnitude of 1% exceedance
 292 probability flows generated by the LOCA CMIP5 ensemble. We see an increase in the difference
 293 between the 5th and 95th percentiles of the ensemble 1% exceedance probability flows for all but one
 294 site, which we address further in the discussion.

295 RCP 4.5 and RCP 8.5 ensemble trends are generally in agreement with one another for each site, with
 296 the RCP 8.5 ensemble typically having a stronger trend compared to the RCP 4.5 ensemble. When
 297 comparing the nonstationary 1% exceedance probability flows to those generated from the stationary
 298 LPIII and GEV distributions, for USGS 11264500 (Merced River), for example, from 1951-2099 generally
 299 between 75%-95% of the nonstationary 1% exceedance probability flows are below the stationary GEV
 300 1% exceedance probability flow. This suggests that while the stationary GEV distribution might generally
 301 have a higher estimate of the seasonal 1% exceedance probability level, there are years where projected

302 covariate conditions would indicate a higher seasonal 1% exceedance probability level with a
 303 nonstationary distribution. When calculating the risk of exceeding some threshold flow over a design
 304 life, if the design life includes a seasonal period where exceedance probability levels are large, the
 305 probability of exceeding that threshold flow will drastically increase. This will be captured in the
 306 simulation results explained further in the results section.

307 Results for USGS 03409500 (Clear Fork River) suggests stationary GEV and LPIII distributions estimate
 308 significantly lower 1% exceedance probability levels compared to the nonstationary model. Results
 309 specific to Clear Fork River in Appendix B show that nonstationary exceedance probability levels diverge
 310 from those of the stationary GEV model for larger exceedance probabilities. There are several extended
 311 periods of missing data for the Clear Fork River site, so limited data could be responsible for diverging
 312 performance between the stationary and nonstationary models for more extreme flows. The stationary
 313 and nonstationary GEV models generate very different 1% exceedance probability levels for USGS
 314 09405500 (North Fork Virgin River). We found the stationary GEV distribution to poorly capture the
 315 more extreme observed historical floods. We also see from the likelihood ratio test there is a great
 316 degree of confidence (Table 1) that the log-likelihood of the nonstationary model is better than that of
 317 the stationary GEV distribution for this site.

318 A spatial plot of the percent change of the median 1% exceedance probability flow generated from the
 319 RCP 8.5 LOCA CMIP5 ensemble between the 1951-1979 and 2040-2069 periods for each site appears in
 320 Figure 5. We see a decrease in the median CMIP 1% seasonal exceedance flows for our study sites in the
 321 southwestern U.S. and an increase in the northern and eastern U.S.



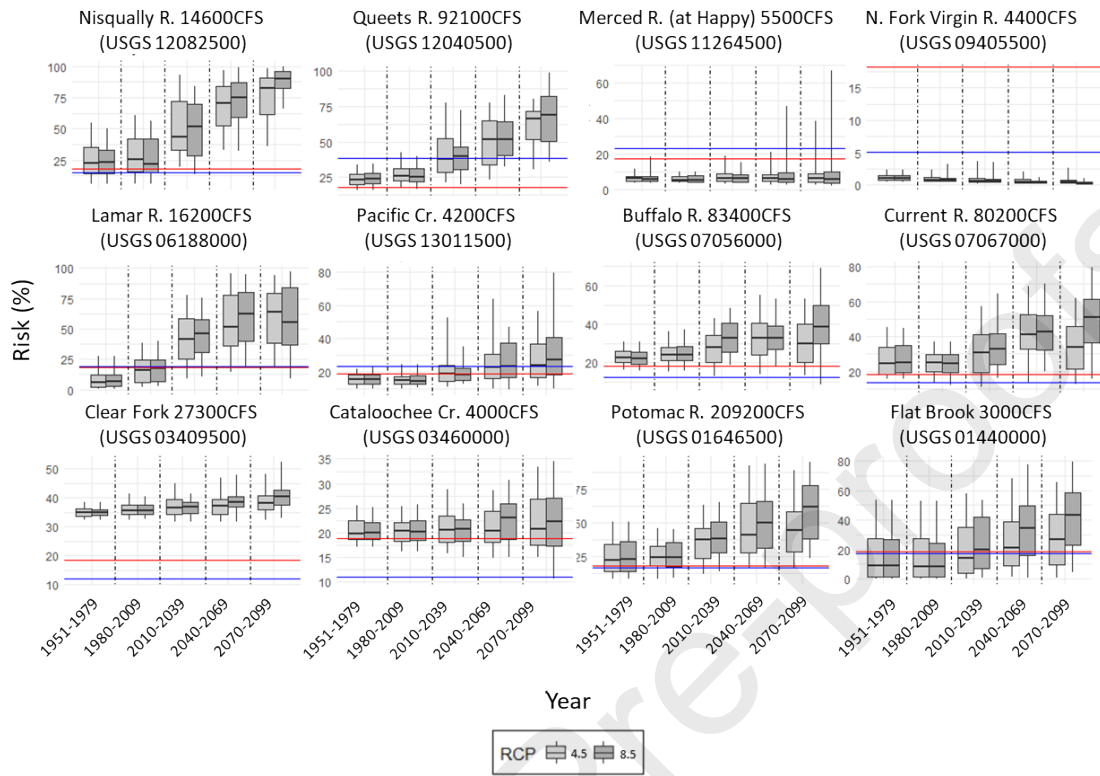
322

323 *Figure 5 Percent change of RCP 8.5 2040-2069 median 1% seasonal exceedance flow compared to the 1951-1979 median 1%*
 324 *exceedance seasonal flow generated from the best GEV model and LOCA CMIP5 covariate projections for each site.*

325 Using the best model and the LOCA CMIP5 ensemble at each site, we simulate the probability of
 326 exceeding a predetermined threshold flow over a specific design life. Here, we select a 20-year design
 327 life using the site's stationary GEV 1% exceedance probability flow as the threshold flow of interest. We
 328 run a large number of simulations for each site and boxplot simulation results. We separate simulations
 329 by the same 30-year periods and by RCP scenario as in Figure 4. Results are shared in Figure 6.
 330 Stationary GEV (red line) and LPIII (blue line) risks are also included. Similar plots for site specific critical
 331 flows are shared in Appendix B.

332

GCM Ensemble Risk of Exceeding Stationary 1% Seasonal Exceedance Probability Flow in 20-year Project



333

334 *Figure 6 Simulation results for the risk of exceeding the site's stationary GEV 1% seasonal exceedance probability flow in a 20-*
 335 *year project life using the best nonstationary GEV model and LOCA CMIP5 covariate projections for each site (boxplots).*
 336 *Stationary LPIII (blue line) and GEV (red line) risks are also included.*

337 While trends between Figure 4 and Figure 6 are similar, we generally see stronger trends in Figure 6.
 338 This is reasonable – a stronger trend will be present when a slight change in seasonal risk is
 339 compounded over 20 years. Further, as we saw in Figure 4, there are years within this LOCA CMIP5
 340 ensemble where covariate conditions result in a much higher seasonal risk compared to the stationary
 341 distribution. If a high-risk season is included in a simulation's 20-year period, the risk of exceeding the
 342 threshold flow over the 20-year period will significantly increase.

343 **5. Discussion and Conclusion**

344 In this paper we utilize the nonstationary generalized extreme value distribution and an ensemble of
 345 climate models to project seasonal 21st century flood risk for twelve sites representing a diverse array of
 346 hydroclimates across the U.S. National Park Service. Results generally project a decrease in seasonal
 347 flood risk for sites in the southwestern U.S. and increases for sites in the eastern and northwestern U.S.
 348 These seasonal results display similar patterns to those identified by Hirsch and Ryberg (2012), who
 349 explored changes in historical flood magnitude under rising carbon dioxide levels at 200 sites across the
 350 U.S. Thus, our projections suggest the trends identified over the time period of Hirsch and Ryberg's
 351 work (where the median record length was 1916-2008) are likely to continue. Further, for many sites
 352 we find flows generated from a hydrologic model improved performance of nonstationary generalized
 353 extreme value distributions when used as covariates.

354 For long-term climate impact studies, two dominant sources of uncertainty arise when using an
355 ensemble of climate models – future scenario uncertainty and model uncertainty (Hawkins and Sutton
356 2009). Described further in Deser et al. (2012), future scenario uncertainty can refer to, for example,
357 uncertainty in greenhouse gas representative concentration pathway trajectories. Our results present
358 only RCP 4.5 and RCP 8.5 scenarios, which typically display similar trends with stronger shifts in flood risk
359 associated with RCP 8.5 trajectories. Model uncertainty arises from the fact that different climate
360 models, given the same forcing, have different responses. As mentioned, for many of our sites, we see
361 increases in the interquartile range and 5th-95th percentile range in our ensemble results presented in
362 Figure 4 and Figure 6. This increase could relate to model uncertainty – climate models with different
363 physical and numerical parameterizations can have diverging responses to long-term projections of
364 input forcings. Our use of 32 GCMs, in part, characterizes this model uncertainty, and one common
365 technique to combine results from climate ensembles involves taking a simple or weighted average of
366 ensemble results (Tebaldi and Knutti 2007). The median and interquartile range in the boxplots
367 presented in Figure 4 and Figure 6 represent this central tendency of the ensemble results, noting that
368 we utilize the same 32 GCMs for each site and we do not assess individual GCM model performance for
369 each site.

370 Our selection of covariates for the best model at each site assumes these general, seasonal average
371 covariates represent the dominant driving mechanisms for seasonal peak flows; shifts in covariates
372 suggest a shift in flood risk due to these dominant flood mechanisms. However, multiple flood
373 generating mechanisms can be present (Berghuijs et al. 2016) and dominant flood mechanisms might
374 exhibit long-term changes (e.g., transitions from snowmelt to rainfall-runoff) (Knowles et al. 2006; Das
375 et al. 2013). For example, covariates like precipitation would not capture flood behavior shifts that arise
376 from more precipitation falling as rainfall (as opposed to snowfall) in the future. This could be an
377 advantage of using VIC model-generated flow as a covariate as we did here - shifting flood generating
378 mechanism behavior for a basin can be better captured because VIC accounts for water and energy
379 balance aspects of a system. In our study, VIC model generated flows are more frequently evident as a
380 best model covariate over precipitation. Shifting dominant flood mechanisms might also shift the timing
381 of peak floods out of the seasons studied in this paper, which cannot be accounted for in our covariates.
382 Additionally, it is important to note the biases that remain in these VIC models (Table A-1) and in the
383 LOCA CMIP5 dataset following bias correction and downscaling processes (Pierce et al. 2014, 2015).

384 We acknowledge this study does not assess nonstationary GEV model parameter uncertainties.
385 Assessing standard errors from maximum likelihood estimates can provide more information on best
386 model performance and exploring this method in a Bayesian framework could also be valuable for
387 assessing uncertainties (Katz et al. 2002; Renard et al. 2013; Cheng et al. 2014; Bracken et al. 2018).
388 Further, investigating the sensitivity of exceedance probability levels generated from the set of models
389 we fit for each site (i.e., the models with different combinations of covariates) could also provide insight
390 into model performance and the relationship between covariates and flood risk that these models
391 capture.

392 Beyond investigating the uncertainties described above, future studies could investigate other flood
393 characteristics like duration. Further, performing this analysis on each season of a year could provide
394 more information on annual peak flood behavior for a particular site.

395 Overall our projected shifts in future flood behavior can help NPS managers assess the need to develop
396 climate change informed flood risk management plans at different park units. This can improve risk

397 mitigation for cultural and natural resources, inform site selection and design for roads, trails, and other
398 infrastructure, and help managers proactively plan for trail and facility closures to ensure visitor safety.
399 Due to the sensitive nature of flood planning for certain projects, we suggest utilizing these results,
400 along with an in-depth understanding of specific basins and other industry accepted flood hazard
401 evaluation techniques, to assess the factor of safety required for flood planning under a changing
402 climate.

403

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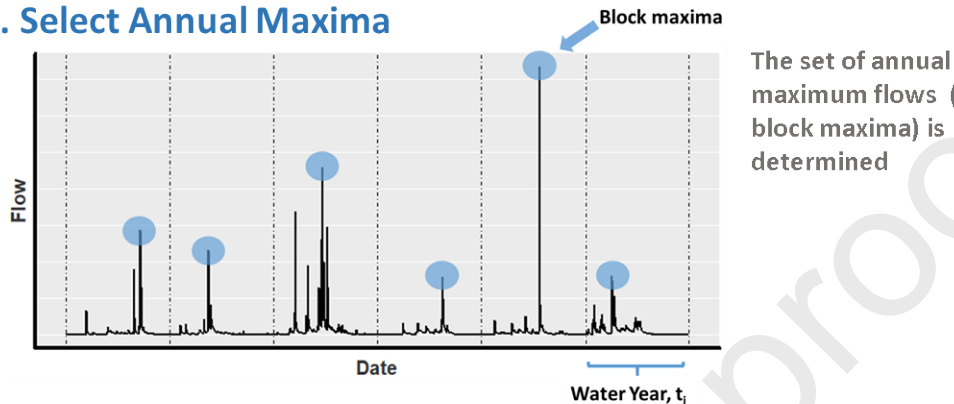
Journal Pre-proofs

Appendix A

The following pages i-iii include a process summary of the stationary and nonstationary flood frequency analysis used in this paper. The process in these pages is demonstrated in an annual timeframe, though it is just as applicable to a seasonal timeframe.

Stationary Flood Risk

1. Select Annual Maxima



2. Fit a GEV Distribution to Block Maxima

A stationary generalized extreme value (GEV) distribution is fit to the block maxima

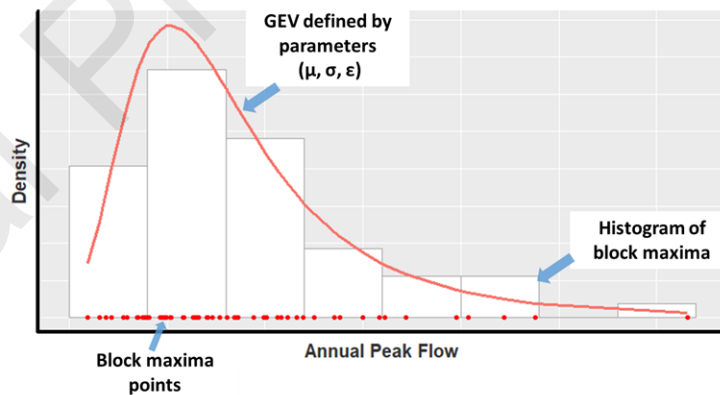
$$F(x) = \exp \left\{ - \left[1 + \epsilon \frac{(x-\mu)}{\sigma} \right]^{-\frac{1}{\epsilon}} \right\}$$

The stationary GEV distribution is defined by 3 parameters:

location (μ), scale (σ), and shape (ϵ)

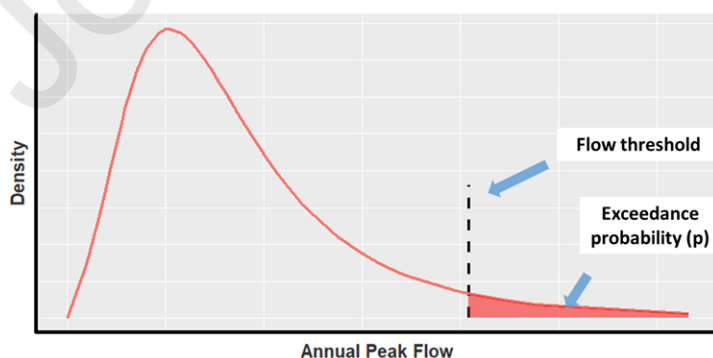
We estimate these parameters using the method of maximum likelihood

μ	ϕ	ϵ
1406.0	994.3	0.6



3. Calculate Risk of Flood

The fit distribution is used to assess the risk of exceeding a flow



$$R = 1 - (1 - p)^n$$

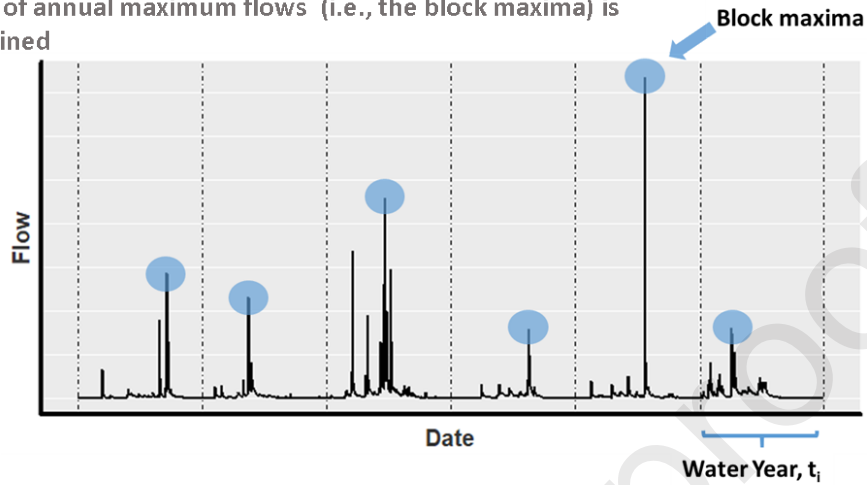
Risk of flood occurring in n years

Probability of flow above flood threshold occurring in any year

Non-Stationary Flood Risk

1. Select Annual Maxima

The set of annual maximum flows (i.e., the block maxima) is determined



2. Fit a Non-Stationary GEV Distribution to Block Maxima

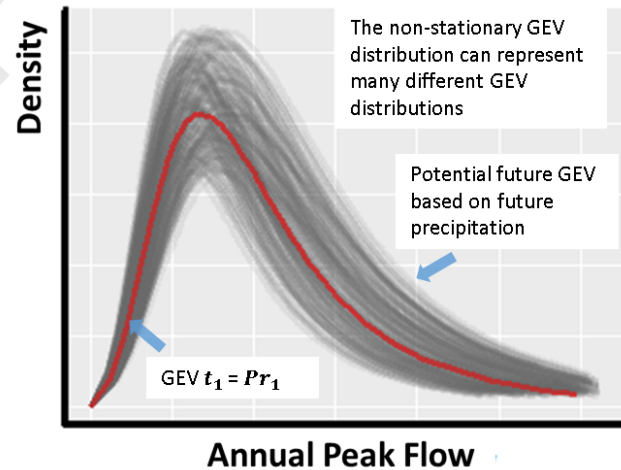
A non-stationary generalized extreme value (GEV) distribution is fit to the block maxima.

A non-stationary GEV distribution allows the location (μ) and scale (σ) parameters to vary with meteorological variables. In this case, the location and scale vary with annual precipitation. However, other climate covariates are often used.

$$F(x) = \exp \left\{ - \left[1 + \varepsilon \frac{(x - \mu(t))}{\sigma(t)} \right]^{-\frac{1}{\varepsilon}} \right\}$$

$$\begin{aligned} \mu(t) &= \beta_0 + \beta_1 Pr_t \\ \sigma(t) &= \alpha_0 + \alpha_1 Pr_t \\ \varepsilon &= \text{constant} \end{aligned}$$

The location and scale shift as a function of precipitation. This GEV distribution is defined by 5 parameters ($\beta_0, \beta_1, \alpha_0, \alpha_1, \varepsilon$)



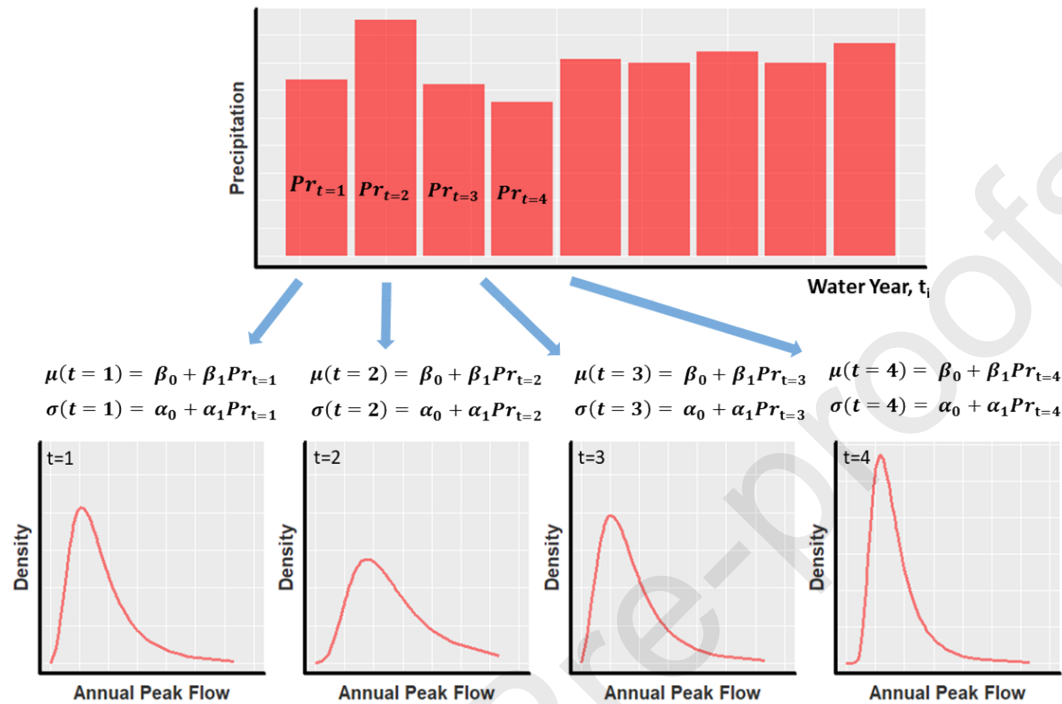
β_0	β_1	α_0	α_1	ε	NIlh	AIC
1324.2	115.8	6.8	0.1	0.2	463.2	936.5

LR Test Against Stationary GEV Model: p-value=0.02

We solve for the 5 parameters using the method of maximum likelihood and assess the model fit using AIC and a likelihood ratio test.

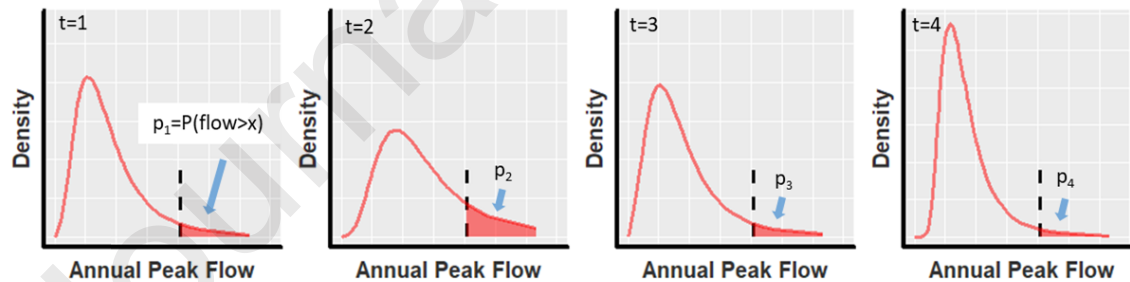
3. Determine the GEV Distribution for Every Time Period

Using the annual precipitation covariate and the fit non-stationary GEV model, a location and scale is determined for each year



4. Calculate Risk by Combining GEVs Calculated for Each Step

The fit distribution is used to assess the risk of exceeding a flow. Because the non-stationary GEV model defines a different location and scale for each year (depending on annual precipitation), the risk of exceeding a flow changes between years.



$$R = 1 - \prod_{t=1}^n (1 - p_t)$$

In this case, the risk calculated is specific to years $t=1$ to $t=4$. Different GEVs, and therefore a different risk, will occur in years $t=5$ to $t=8$, for example.

	USGS Site	USGS Site Description	Season	VIC Model Skill Scores		
				NSE	COR	PBIAS (%)
1	12082500	Nisqually River near National, WA	Oct-Mar	-0.19	0.67	-42.0
			Oct-Dec	-0.05	0.79	-45.9
2	12040500	Queets River near Clearwater, WA	Oct-Mar	0.62	0.92	-23.9
			Oct-Dec	0.65	0.93	-27.7
3	11264500	Merced River at Happy Isles Bridge near Yosemite, CA	Apr-Jun	-0.42	0.78	19.3
			Nov-Mar	-0.05	0.69	-72.1
4	09405500	North Fork Virgin River near Springdale, UT	Apr-Jun	0.34	0.79	-1.7
			Nov-Mar	-1.30	0.66	-11.8
5	06188000	Lamar River near Tower Ranger Station, YNP	Apr-Jun	-0.46	0.77	-69.4
			Nov-Mar	-1.20	0.53	-4.1
6	13011500	Pacific Creek at Moran, WY	Apr-Jun	0.26	0.82	-48.6
			Nov-Mar	-0.16	0.61	-19.4
7	07056000	Buffalo River near St. Joe, AR	Jan-May	0.77	0.89	-14.1
			Oct-Dec	0.87	0.95	21.2
8	07067000	Current River at Van Buren, MO	Feb-May	0.55	0.82	-9.8
			Nov-Jan	0.55	0.92	20.6
9	03409500	Clear Fork near Robbins, TN	Dec-Mar	0.53	0.82	-23.4
			Oct-Nov	0.37	0.85	-28.5
10	03460000	Cataloochee Creek near Cataloochee, NC	Dec-Mar	0.18	0.87	19.8
			Oct-Nov	0.60	0.93	-6.0
11	01646500	Potomac River near Washington, D.C. Little Falls Pump	Feb-May	0.71	0.85	-6.0
			Nov-Jan	0.59	0.84	27.3
12	01440000	Flat Brook near Flatbrookville, NJ	Feb-Apr	0.43	0.78	-20.6
			Nov-Jan	0.64	0.82	-10.2

Table A-1 Monthly VIC model flow performance metrics over seasons of interest (using observed historical forcings from 1951-2005 water years) for each study site. Metrics include the Nash-Sutcliffe efficiency, Pearson correlation, and percent bias.

	USGS Site	USGS Site Description	Season	COR
1	12082500	Nisqually River near National, WA	Oct-Mar	0.68
			Oct-Dec	0.59
2	12040500	Queets River near Clearwater, WA	Oct-Mar	0.62
			Oct-Dec	0.50
3	11264500	Merced River at Happy Isles Bridge near Yosemite, CA	Apr-Jun	0.72
			Nov-Mar	0.14
4	09405500	North Fork Virgin River near Springdale, UT	Apr-Jun	0.80
			Nov-Mar	0.55
5	06188000	Lamar River near Tower Ranger Station, YNP	Apr-Jun	0.52
			Nov-Mar	-0.09
6	13011500	Pacific Creek at Moran, WY	Apr-Jun	0.69
			Nov-Mar	0.12
7	07056000	Buffalo River near St. Joe, AR	Jan-May	0.61
			Oct-Dec	-0.05
8	07067000	Current River at Van Buren, MO	Feb-May	0.60
			Nov-Jan	0.26
9	03409500	Clear Fork near Robbins, TN	Dec-Mar	0.59
			Oct-Nov	-0.03
10	03460000	Cataloochee Creek near Cataloochee, NC	Dec-Mar	0.50
			Oct-Nov	0.32
11	01646500	Potomac River near Washington, D.C. Little Falls Pump	Feb-May	0.70
			Nov-Jan	0.30
12	01440000	Flat Brook near Flatbrookville, NJ	Feb-Apr	0.65
			Nov-Jan	0.03

Table A-2 Correlations between season of interest and previous season of interest average mean daily VIC model flow and the site's season of interest peak mean daily flow. For each site, the top set of months correspond to the season of interest and the bottom set of months corresponding to the previous season of interest.

Modeling Center (or Group)	Institute ID	Model Name
Commonwealth Scientific and Industrial Research Organization (CSIRO) and Bureau of Meteorology (BOM), Australia	CSIRO-BOM	ACCESS1.0
		ACCESS1.3
Beijing Climate Center, China Meteorological Administration	BCC	BCC-CSM1.1
		BCC-CSM1.1(m)
Canadian Centre for Climate Modelling and Analysis National Center for Atmospheric Research	CCCMA	CanESM2
	NCAR	CCSM4
Community Earth System Model Contributors	NSF-DOE-NCAR	CESM1(BGC)
		CESM1(CAM5)
Centro Euro-Mediterraneo per i Cambiamenti Climatici	CMCC	CMCC-CM
		CMCC-CMS
Centre National de Recherches Météorologiques / Centre Européen de Recherche et Formation Avancée en Calcul Scientifique	CNRM-CERFACS	CNRM-CM5
Commonwealth Scientific and Industrial Research Organization in collaboration with Queensland Climate Change Centre of Excellence	CSIRO-QCCCE	CSIRO-Mk3.6.0
EC-EARTH consortium	EC-EARTH	EC-EARTH
LASG, Institute of Atmospheric Physics, Chinese Academy of Sciences and CESS, Tsinghua University	LASG-CESS	FGOALS-g2
NOAA Geophysical Fluid Dynamics Laboratory	NOAA GFDL	GFDL-CM3
		GFDL-ESM2G
NASA Goddard Institute for Space Studies	NASA GISS	GFDL-ESM2M
		GISS-E2-H
National Institute of Meteorological Research/Korea Meteorological Administration	NIMR/KMA	GISS-E2-R
		HadGEM2-AO
Met Office Hadley Centre (additional HadGEM2-ES realizations contributed by Instituto Nacional de Pesquisas Espaciais)	MOHC (additional realizations by INPE)	HadGEM2-CC
		HadGEM2-ES
Institute for Numerical Mathematics	INM	INM-CM4
Institut Pierre-Simon Laplace	IPSL	IPSL-CM5A-LR
		IPSL-CM5A-MR
Japan Agency for Marine-Earth Science and Technology, Atmosphere and Ocean Research Institute (The University of Tokyo), and National Institute for Environmental Studies	MIROC	MIROC-ESM
		MIROC-ESM-CHEM
Atmosphere and Ocean Research Institute (The University of Tokyo), National Institute for Environmental Studies, and Japan Agency for Marine-Earth Science and Technology	MIROC	MIROC5
Max-Planck-Institut für Meteorologie (Max Planck Institute for Meteorology)	MPI-M	MPI-ESM-MR
		MPI-ESM-LR
Meteorological Research Institute	MRI	MRI-CGCM3
Norwegian Climate Centre	NCC	NorESM1-M

Table A-3 The 32 CMIP5 models in the LOCA CMIP5 ensemble. We acknowledge the World Climate Research Program's Working Group on Coupled Modelling, which is responsible for CMIP, and we thank the climate modeling groups for producing and making available their model output. For CMIP the U.S. Department of Energy's Program for Climate Model Diagnosis and Intercomparison provides coordinating support and led development of software infrastructure in partnership with the Global Organization for Earth System Science Portals.

Appendix B**Seasonal Flow Analysis**

The following pages contain seasonal flood frequency results for all 12 sites. Sites and the season of interest are shared in the table below.

	USGS Site	USGS Site Description	Unit	Season of Analysis	Previous Season Covariates
1	12082500	Nisqually River near National, WA	Mount Rainier National Park	Oct-Mar	Oct-Dec
2	12040500	Queets River near Clearwater, WA	Olympic National Park	Oct-Mar	Oct-Dec
3	11264500	Merced River at Happy Isles Bridge near Yosemite, CA	Yosemite National Park	Apr-Jun	Nov-Mar
4	09405500	North Fork Virgin River near Springdale, UT	Zion National Park	Apr-Jun	Nov-Mar
5	06188000	Lamar River near Tower Ranger Station, YNP	Yellowstone National Park	Apr-Jun	Nov-Mar
6	13011500	Pacific Creek at Moran, WY	Grand Teton National Park	Apr-Jun	Nov-Mar
7	07056000	Buffalo River near St. Joe, AR	Buffalo National River	Jan-May	Oct-Dec
8	07067000	Current River at Van Buren, MO	Ozark National Scenic Riverways	Feb-May	Nov-Jan
9	03409500	Clear Fork near Robbins, TN	Big South Fork National River and Rec. Area	Dec-Mar	Oct-Nov
10	03460000	Cataloochee Creek near Cataloochee, NC	Great Smoky Mountains National Park	Dec-Mar	Oct-Nov
11	01646500	Potomac River near Washington, D.C. Little Falls Pump	Chesapeake & Ohio Canal National Hist. Park	Feb-May	Nov-Jan
12	01440000	Flat Brook near Flatbrookville, NJ	Delaware Water Gap National Rec. Area	Feb-Apr	Nov-Jan

Seasonal Extreme Flow Analysis
 Nisqually River near National, WA (USGS 12082500)
 Mount Rainier National Park

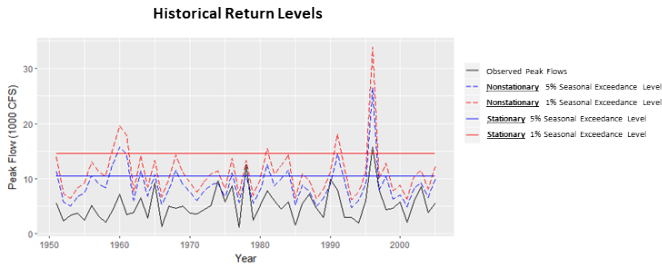


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Best Nonstationary GEV Model
 Location ~ VIC, Scale ~ VIC

Location	Scale		Shape	NIlh	AIC	BIC
μ_0	μ_1	ϕ_0	ϕ_1			
-0.35	10.33	-0.46	1.998	-0.05	109.7	229.5

LR Test Against Stationary GEV Model: **p-value=3.842e-08**

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

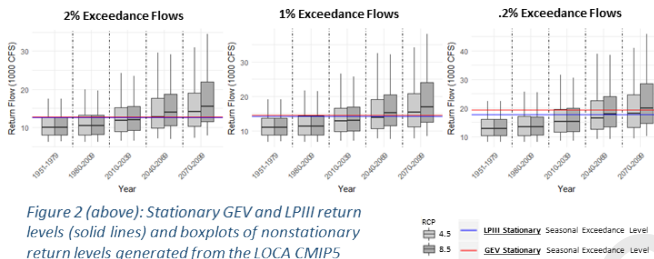


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

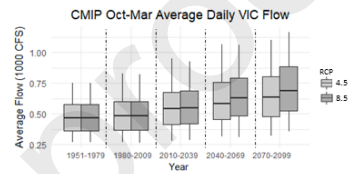


Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	RCP			Nonstationary GEV Model - GCM Ensemble Percentile						
			25th	Median	75th	25th	Median	75th	25th	Median	75th	
20%	7.0	7.2	1980-2009			2040-2069			2070-2099			
			4.5	5.3	7.0	8.9	6.6	8.6	11.6	6.9	9.5	12.5
10%	8.7	8.9	8.5	5.4	7.0	8.9	7.1	9.4	12.3	7.7	10.3	14.2
			4.5	6.2	8.1	10.3	7.6	10.0	13.6	8.0	11.0	14.7
4%	11.0	11.0	8.5	6.3	8.2	10.3	8.2	10.9	14.4	8.9	12.0	16.7
			4.5	7.3	9.5	12.0	8.9	11.6	15.9	9.4	12.8	17.2
1%	14.6	14.2	8.5	7.3	9.5	12.0	9.5	12.7	16.9	10.4	14.1	19.7
			4.5	8.8	11.4	14.4	10.7	13.9	19.2	11.2	15.4	20.8
			8.5	8.8	11.4	14.4	11.4	15.2	20.4	12.5	16.9	23.9

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

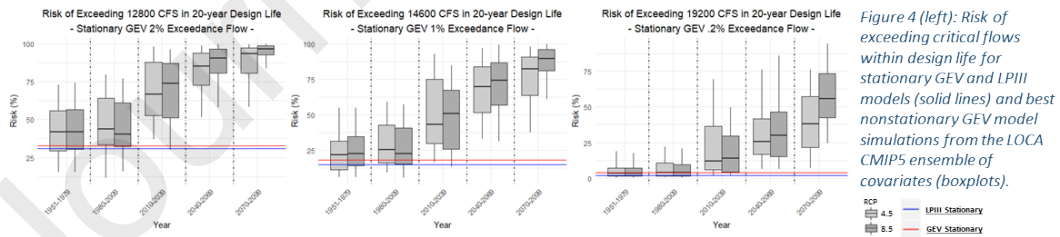


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis Queets River near Clearwater, WA (USGS 12040500) Olympic National Park

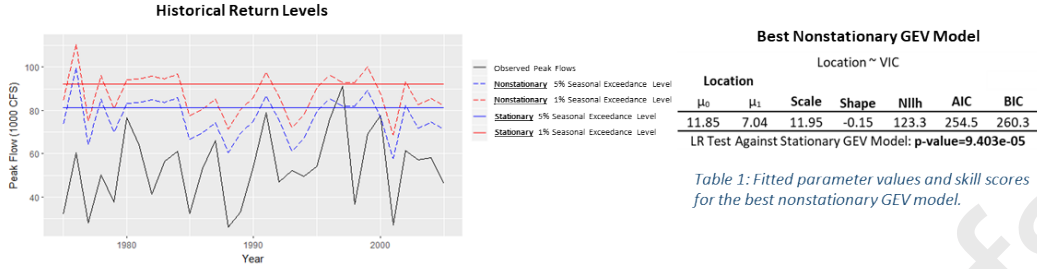


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

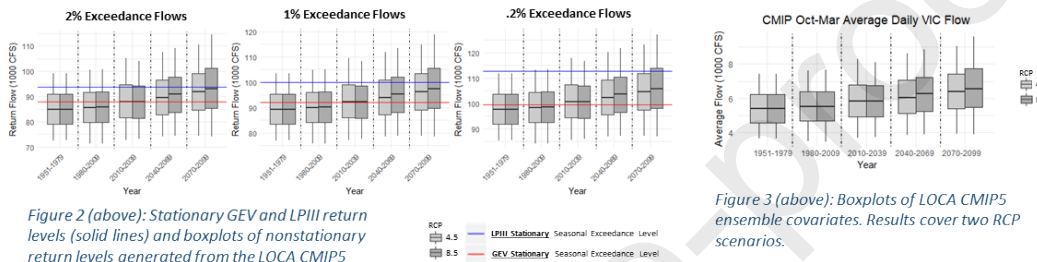


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	RCP	Nonstationary GEV Model - GCM Ensemble Percentile			2040-2069			2070-2099		
				25th	Median	75th	25th	Median	75th	25th	Median	75th
20%	67.4	67.7	4.5	60.7	66.6	72.7	63.8	70.4	77.6	65.7	72.7	80.0
			8.5	60.8	66.8	72.7	64.6	71.9	78.7	66.4	74.0	82.2
10%	75.1	76.8	4.5	67.5	73.4	79.4	70.5	77.1	84.3	72.5	79.5	86.8
			8.5	67.5	73.5	79.4	71.4	78.6	85.4	73.1	80.7	88.9
4%	83.1	87.0	4.5	74.9	80.8	86.9	77.9	84.6	91.8	79.9	86.9	94.2
			8.5	74.9	81.0	86.9	78.8	86.1	92.9	80.6	88.2	96.4
1%	92.1	100.0	4.5	84.1	90.0	96.1	87.1	93.8	101.0	89.1	96.1	103.4
			8.5	84.1	90.2	96.1	88.0	95.3	102.1	89.8	97.4	105.6

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

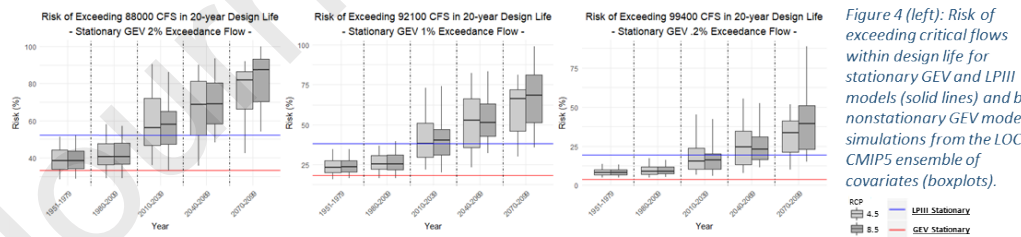


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis Merced River at Happy Isles Bridge near Yosemite, CA (USGS 11264500) Yosemite National Park

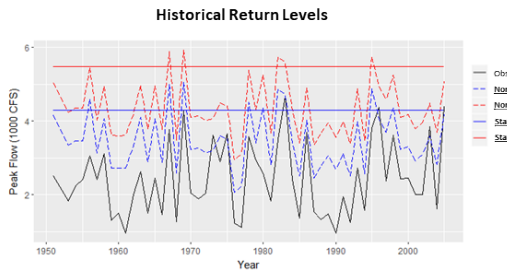


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

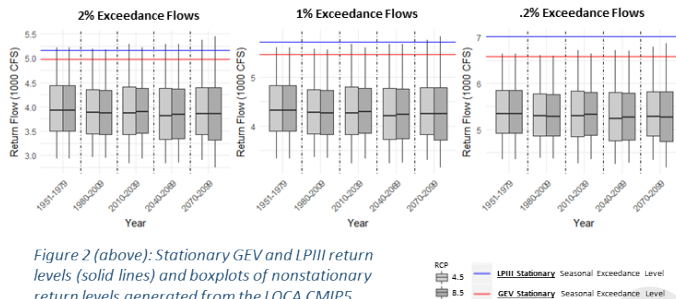


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

Best Nonstationary GEV Model
Location \sim Pre_Pr + Pr

Location	μ_0	μ_1	μ_2	Scale	Shape	NIH	AIC	BIC
	0.11	0.24	0.29	0.37	0.10	35.0	79.9	89.9

LR Test Against Stationary GEV Model: $p\text{-value} < 2.2e-16$

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

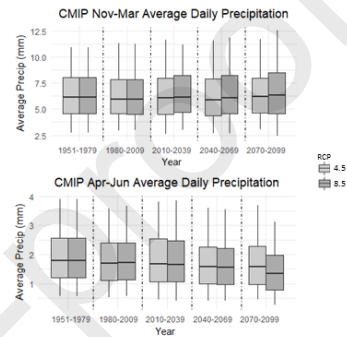


Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models			Nonstationary GEV Model - GCM Ensemble Percentile								
	GEV	LPIII	RCP	1980-2009			2040-2069			2070-2099		
				25th	Median	75th	25th	Median	75th	25th	Median	75th
20%	3.2	3.2	4.5	2.3	2.7	3.2	2.2	2.6	3.2	2.3	2.7	3.2
8.5			2.3	2.7	3.2	2.2	2.7	3.2	2.2	2.7	3.2	
10%	3.8	3.8	4.5	2.6	3.0	3.5	2.5	3.0	3.6	2.6	3.0	3.6
8.5			2.6	3.0	3.5	2.5	3.0	3.5	2.5	3.0	3.6	
4%	4.5	4.6	4.5	3.1	3.5	4.0	3.0	3.4	4.0	3.1	3.5	4.0
8.5			3.1	3.5	4.0	3.0	3.5	4.0	2.9	3.5	4.0	
1%	5.5	5.7	4.5	3.8	4.3	4.7	3.7	4.2	4.8	3.8	4.2	4.8
8.5			3.8	4.3	4.7	3.7	4.2	4.8	3.7	4.2	4.8	

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

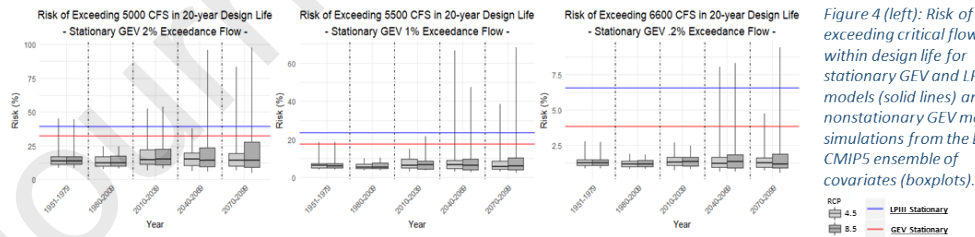


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis North Fork Virgin River near Springdale, UT (USGS 09405500) Zion National Park

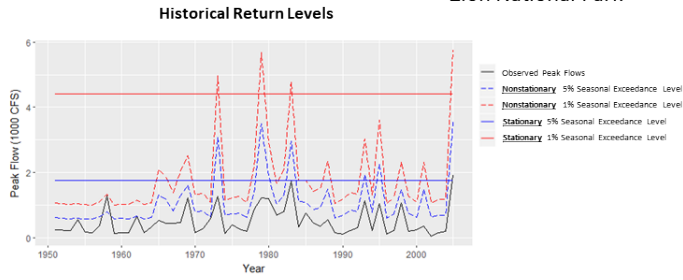


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Best Nonstationary GEV Model

Location ~ VIC, Scale ~ VIC

Location		Scale			Shape	Nllh	AIC	BIC
μ_0	μ_1	ϕ_0	ϕ_1					
0.13	1.25	-2.36	1.740	0.27	-13.7	-17.4	-7.3	

LR Test Against Stationary GEV Model: **p-value=1.621e-13**

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

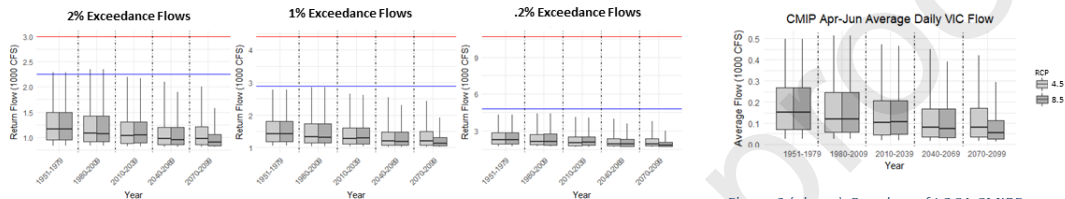


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	Nonstationary GEV Model - GCM Ensemble Percentile									
			RCP	1980-2009			2040-2069			2070-2099		
				25th	Median	75th	25th	Median	75th	25th	Median	75th
20%	0.7	0.8	4.5	0.4	0.5	0.7	0.4	0.4	0.6	0.4	0.4	0.6
			8.5	0.4	0.5	0.7	0.4	0.4	0.6	0.3	0.4	0.5
10%	1.2	1.1	4.5	0.5	0.6	0.9	0.5	0.6	0.7	0.5	0.6	0.7
			8.5	0.5	0.6	0.9	0.5	0.5	0.7	0.5	0.5	0.6
4%	2.0	1.7	4.5	0.7	0.9	1.2	0.7	0.8	1.0	0.7	0.8	1.0
			8.5	0.7	0.9	1.2	0.7	0.8	1.0	0.7	0.7	0.8
1%	4.4	2.9	4.5	1.1	1.3	1.7	1.1	1.2	1.5	1.1	1.2	1.5
			8.5	1.1	1.3	1.7	1.1	1.2	1.5	1.1	1.1	1.3

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

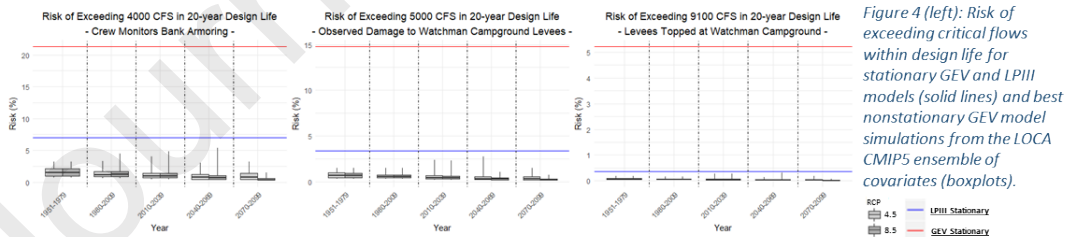


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis Lamar River near Tower Ranger Station, YNP (USGS 06188000) Yellowstone National Park

Historical Return Levels

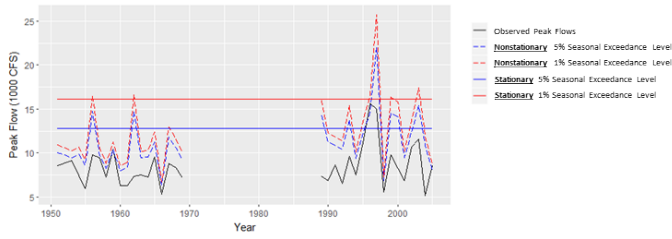


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Best Nonstationary GEV Model

Location ~ Pre_VIC+VIC, Scale ~ Pre_VIC+VIC

Location		Scale					Shape	Nllh	AIC	BIC
μ_0	μ_1	μ_2	ϕ_0	ϕ_1	ϕ_2					
7.04	-19.72	3.76	-0.06	-7.95	1.52	-0.20	62.1	138.2	149.3	

LR Test Against Stationary GEV Model: p-value=8.761e-06

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

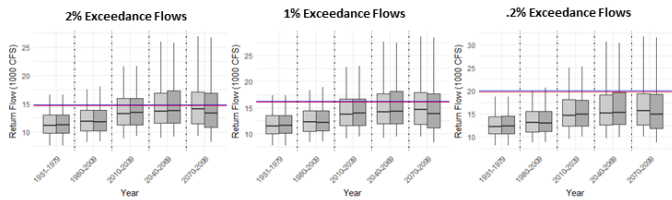


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

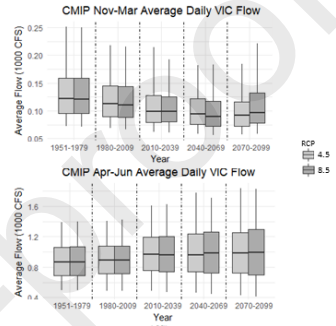


Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	RCP	Nonstationary GEV Model - GCM Ensemble Percentile			2040-2069			2070-2099		
				25th	Median	75th	25th	Median	75th	25th	Median	75th
20%	10.0	10.1	4.5	8.7	9.9	11.2	9.6	11.0	13.1	9.6	11.3	13.2
				8.5	8.7	9.8	11.2	9.7	11.1	13.3	9.1	10.9
10%	11.4	11.5	4.5	9.2	10.6	12.2	10.3	12.0	14.4	10.3	12.3	14.6
				8.5	9.3	10.5	12.2	10.3	12.1	14.8	9.8	11.8
4%	13.3	13.3	4.5	9.8	11.4	13.2	11.0	13.0	15.9	11.0	13.4	16.1
				8.5	9.9	11.3	13.2	11.1	13.1	16.3	10.4	12.8
1%	16.2	16.3	4.5	10.5	12.3	14.4	11.9	14.2	17.7	11.9	14.6	17.9
				8.5	10.6	12.2	14.4	12.0	14.3	18.1	11.2	13.9

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

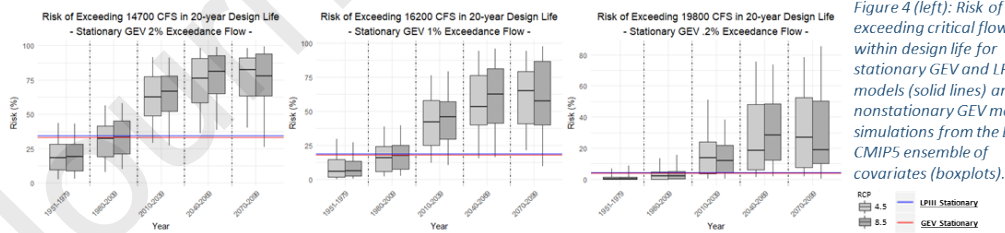


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis
 Pacific Creek at Moran, WY (USGS 13011500)
 Grand Teton National Park

Historical Return Levels

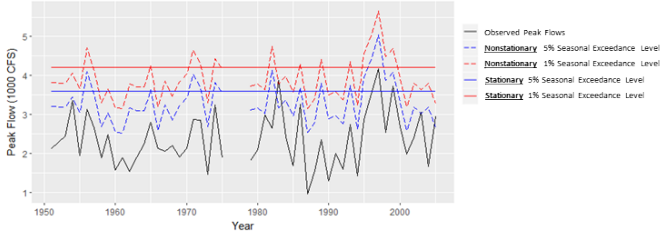


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Best Nonstationary GEV Model

Location ~ Pre_Pr + Pr

Location						
μ_0	μ_1	μ_2	Scale	Shape	NIH	BIC
-0.17	0.52	0.29	0.38	-0.004	31.9	83.5

LR Test Against Stationary GEV Model: p-value = 4.235e-10

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

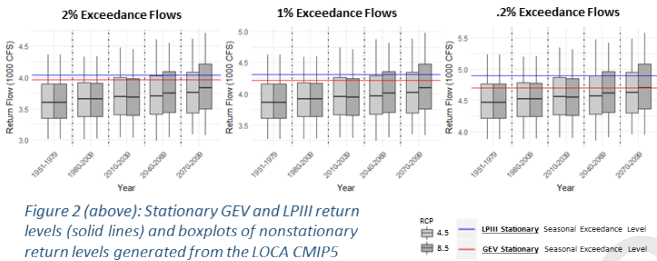


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

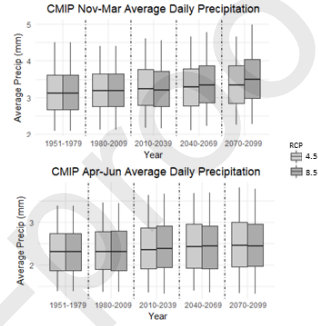


Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)										
	GEV	LPIII	Nonstationary GEV Model - GCM Ensemble Percentile			2040-2069			2070-2099				
			RCP	25th	Median	75th	25th	Median	75th	25th	Median	75th	
20%	2.9	2.9	1980-2009	4.5	2.5	2.7	3.0	2.5	2.8	3.1	2.5	2.8	3.2
			8.5	2.5	2.7	3.0	2.5	2.8	3.2	2.6	2.9	3.3	
10%	3.3	3.3	1980-2009	4.5	2.8	3.0	3.3	2.8	3.1	3.4	2.8	3.1	3.5
			8.5	2.7	3.0	3.3	2.8	3.1	3.5	2.9	3.2	3.6	
4%	3.7	3.7	1980-2009	4.5	3.1	3.4	3.7	3.1	3.4	3.8	3.2	3.5	3.8
			8.5	3.1	3.4	3.6	3.2	3.5	3.8	3.2	3.6	3.9	
1%	4.2	4.3	1980-2009	4.5	3.6	3.9	4.2	3.7	4.0	4.3	3.7	4.0	4.3
			8.5	3.6	3.9	4.2	3.7	4.0	4.4	3.8	4.1	4.5	

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

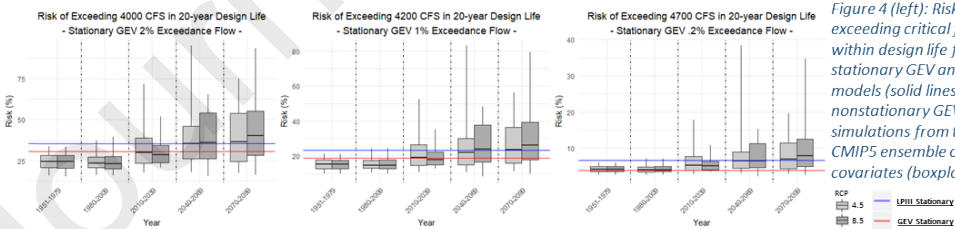


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis Buffalo River near St. Joe, AR (USGS 07056000) Buffalo National River

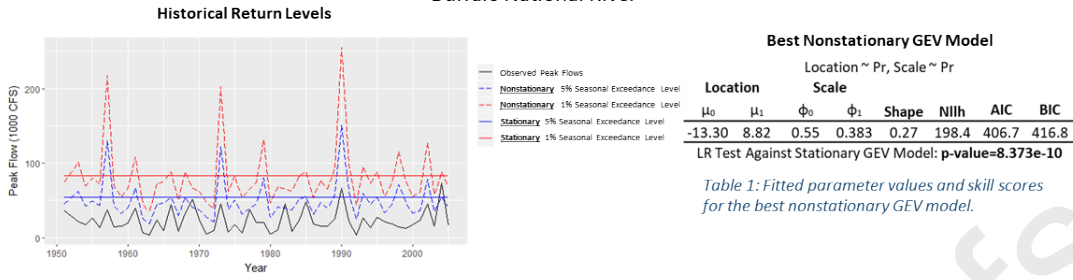


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Best Nonstationary GEV Model

Location ~ Pr, Scale ~ Pr

Location		Scale			Shape	Nilh	AIC	BIC
μ_0	μ_1	ϕ_0	ϕ_1					
-13.30	8.82	0.55	0.383	0.27	198.4	406.7	416.8	

LR Test Against Stationary GEV Model: **p-value=8.373e-10**

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

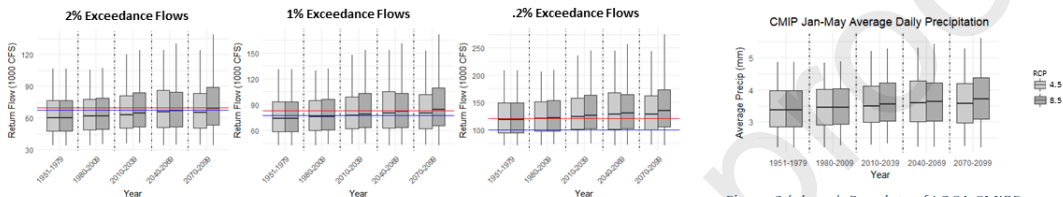


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)										
	GEV	LPIII	Nonstationary GEV Model - GCM Ensemble Percentile										
			RCP	1980-2009			2040-2069			2070-2099			
20%	32.8	34.0	4.5	25th	22.1	29.0	36.9	23.3	31.0	40.9	22.9	30.9	39.6
			8.5	22.3	29.2	37.4	23.5	31.7	40.1	24.5	32.6	42.3	
10%	42.9	44.3	4.5	28.7	37.1	46.8	30.1	39.5	52.0	29.6	39.4	50.4	
			8.5	28.9	37.2	47.6	30.3	40.4	51.0	31.5	41.6	53.8	
4%	57.5	57.6	4.5	39.1	49.9	62.7	40.9	53.1	69.6	40.3	52.9	67.4	
			8.5	39.4	50.1	63.7	41.2	54.2	68.3	42.7	55.7	72.0	
1%	83.4	77.3	4.5	60.4	75.9	95.0	63.0	80.6	105.4	62.0	80.3	102.1	
			8.5	60.8	76.2	96.5	63.4	82.3	103.4	65.5	84.6	109.1	

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

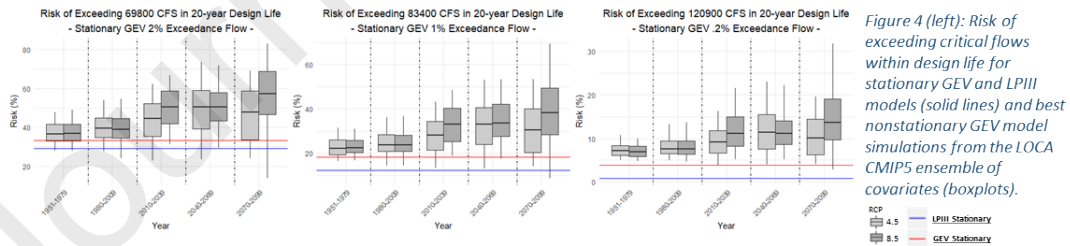


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis Current River at Van Buren, MO (USGS 07067000) Ozark National Scenic Riverways

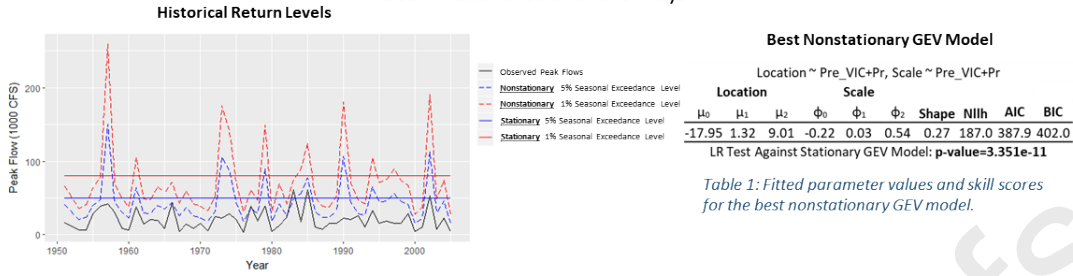


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Best Nonstationary GEV Model

Location ~ Pre_VIC+Pr, Scale ~ Pre_VIC+Pr

Location		Scale					Shape	Nllh	AIC	BIC
μ_0	μ_1	μ_2	Φ_0	Φ_1	Φ_2					
-17.95	1.32	9.01	-0.22	0.03	0.54	0.27	187.0	387.9	402.0	

LR Test Against Stationary GEV Model: **p-value=3.351e-11**

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

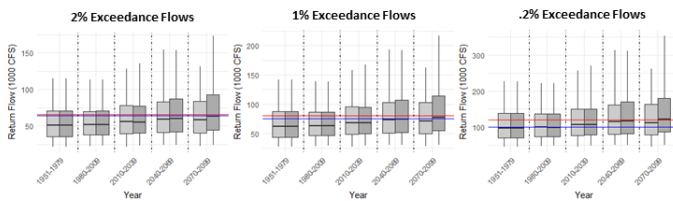


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

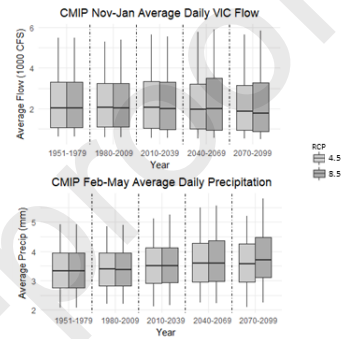


Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	Nonstationary GEV Model - GCM Ensemble Percentile			2040-2069			2070-2099			
			RCP	25th	Median	75th	25th	Median	75th	25th	Median	75th
20%	29.3	30.3	4.5	17.8	25.3	34.2	19.4	28.7	39.6	19.2	28.2	40.1
			8.5	18.0	25.2	34.2	20.1	29.4	41.3	21.1	30.8	43.8
10%	39.0	40.3	4.5	22.6	31.8	43.1	24.6	36.3	50.4	24.5	35.4	50.8
			8.5	22.7	31.7	43.1	25.5	36.9	52.2	26.8	38.7	55.7
4%	53.5	53.8	4.5	30.4	42.1	57.2	33.1	48.4	67.4	32.9	47.2	68.0
			8.5	30.5	41.9	57.3	33.9	49.0	70.1	36.0	51.5	75.0
1%	80.2	75.1	4.5	46.3	63.5	86.2	50.4	72.8	102.6	49.7	70.9	102.9
			8.5	46.8	63.0	86.4	51.4	73.9	106.9	54.6	77.4	113.7

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

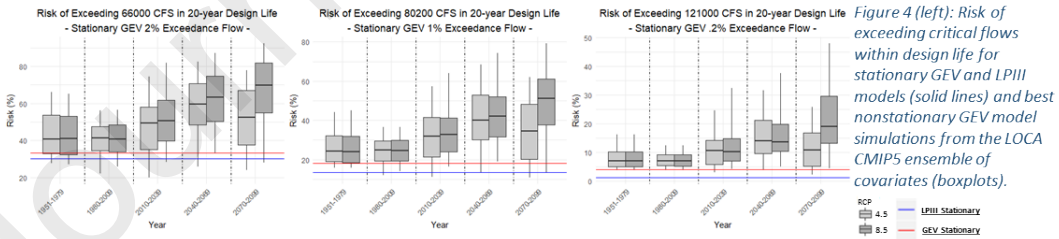


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis
 Clear Fork near Robbins, TN (USGS 03409500)
 Big South Fork National River and Recreation Area

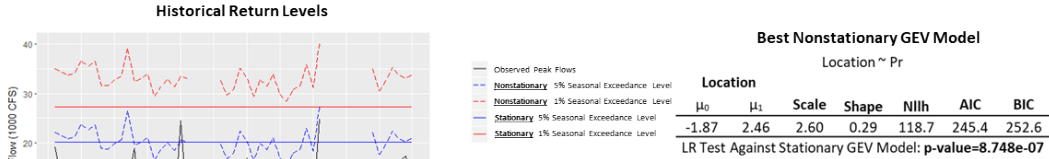


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

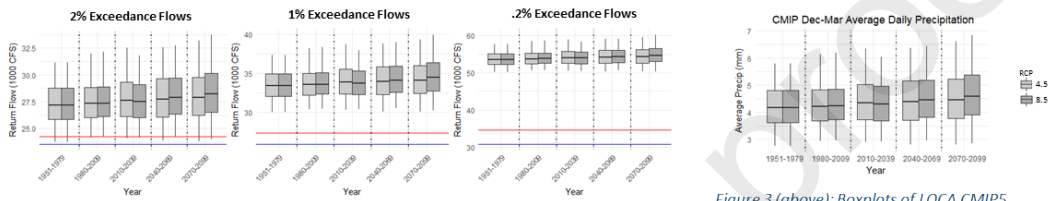


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	Nonstationary GEV Model - GCM Ensemble Percentile									
			RCP	1980-2009			2040-2069			2070-2099		
20%	14.1	14.4	4.5	12.1	13.4	14.9	12.1	13.8	15.7	12.3	14.0	15.9
			8.5	12.2	13.4	14.9	12.4	14.0	15.8	12.6	14.3	16.2
10%	17.2	17.5	4.5	15.4	16.8	18.2	15.5	17.2	19.1	15.7	17.3	19.2
			8.5	15.6	16.8	18.3	15.8	17.3	19.1	16.0	17.7	19.6
4%	21.2	21.0	4.5	20.9	22.2	23.7	20.9	22.6	24.5	21.1	22.8	24.7
			8.5	21.0	22.2	23.7	21.2	22.8	24.6	21.4	23.1	25.1
1%	27.3	25.8	4.5	32.2	33.5	35.0	32.3	34.0	35.8	32.4	34.1	36.0
			8.5	32.3	33.5	35.1	32.5	34.1	35.9	32.7	34.4	36.4

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

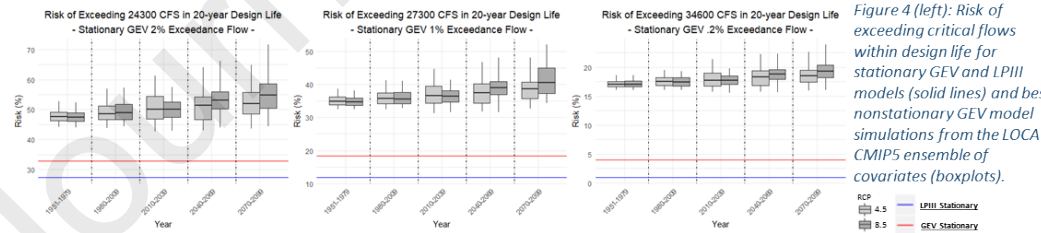


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis
 Cataloochee Creek near Cataloochee, NC (USGS 03460000)
 Great Smoky Mountains National Park

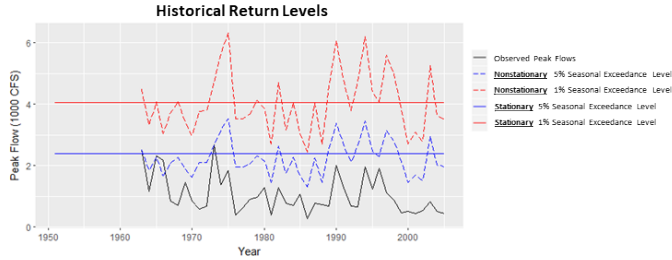


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Best Nonstationary GEV Model

Location ~ VIC, Scale ~ VIC

Location		Scale		Shape	NIrh	AIC	BIC
μ_0	μ_1	ϕ_0	ϕ_1				
0.14	3.12	-1.94	3.175	0.36	21.7	53.4	62.5

LR Test Against Stationary GEV Model: **p-value=2.034e-05**

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

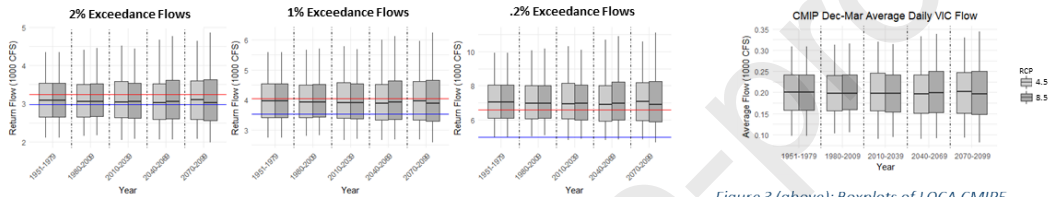


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	Nonstationary GEV Model - GCM Ensemble Percentile									
			RCP	1980-2009			2040-2069			2070-2099		
20%	1.4	1.4	4.5	1.1	1.3	1.5	1.1	1.3	1.5	1.1	1.3	1.5
			8.5	1.1	1.3	1.5	1.1	1.3	1.5	1.1	1.3	1.6
10%	1.9	1.9	4.5	1.4	1.7	2.0	1.4	1.7	2.0	1.4	1.7	2.0
			8.5	1.5	1.7	2.0	1.4	1.7	2.0	1.4	1.7	2.0
4%	2.6	2.5	4.5	2.0	2.4	2.7	2.0	2.4	2.7	2.0	2.4	2.8
			8.5	2.1	2.4	2.7	2.0	2.4	2.8	2.0	2.4	2.8
1%	4.0	3.5	4.5	3.4	3.9	4.5	3.3	3.9	4.5	3.3	4.0	4.6
			8.5	3.4	3.9	4.5	3.4	3.9	4.6	3.3	3.9	4.6

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

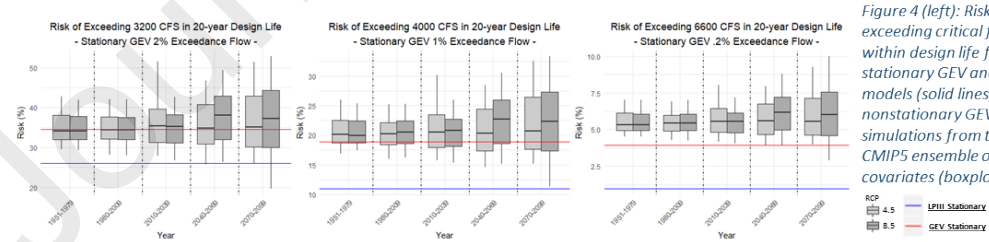


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis

Potomac River near Washington, D.C. Little Falls Pump (USGS 01646500)
Chesapeake & Ohio Canal National Historical Park

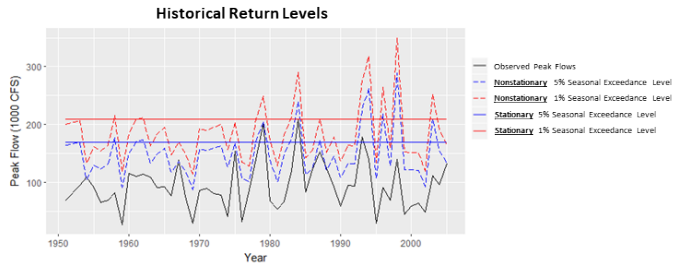


Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Best Nonstationary GEV Model

Location ~ VIC, Scale ~ VIC

Location	Scale	μ_0	μ_1	ϕ_0	ϕ_1	Shape	NIllh	AIC	BIC
-1.87	4.58	2.60	0.029	-0.02	258.7	527.4	537.5		

LR Test Against Stationary GEV Model: **p-value=8.749e-10**

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

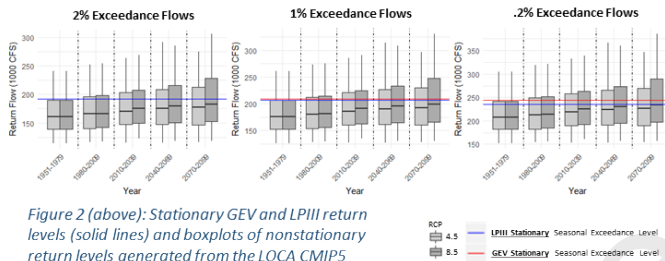


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

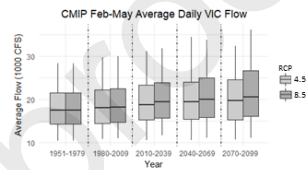


Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	RCP	Nonstationary GEV Model - GCM Ensemble Percentile								
				1980-2009			2040-2069			2070-2099		
				25th	Median	75th	25th	Median	75th	25th	Median	75th
20%	127.0	129.1	4.5	94.2	114.5	138.0	100.2	121.9	148.2	99.2	123.5	151.0
			8.5	96.0	115.0	139.6	102.4	125.5	153.5	103.9	128.2	163.2
10%	149.1	151.4	4.5	108.9	130.8	156.4	115.3	138.9	167.6	114.2	140.6	170.7
			8.5	110.8	131.4	158.2	117.7	142.8	173.4	119.4	145.7	184.1
4%	174.9	176.2	4.5	127.1	151.1	179.3	134.1	160.0	191.7	132.9	161.8	195.1
			8.5	129.2	151.7	181.2	136.7	164.3	198.2	138.5	167.4	210.0
1%	209.2	206.9	4.5	153.4	180.2	212.2	161.2	190.3	226.3	159.8	192.3	230.2
			8.5	155.7	180.9	214.4	164.1	195.1	233.7	166.1	198.7	247.4

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

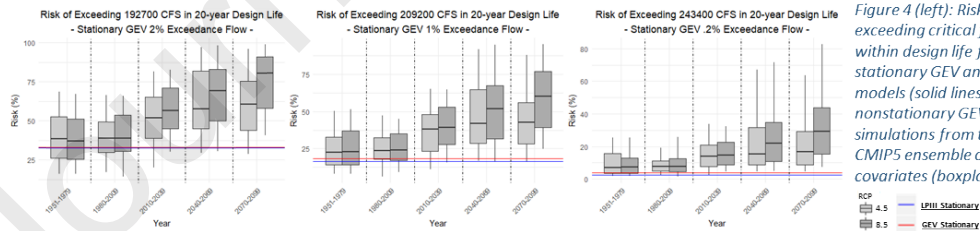


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Seasonal Extreme Flow Analysis
 Flat Brook near Flatbrookville, NJ (USGS 01440000)
 Delaware Water Gap National Recreation Area



Figure 1 (above): A comparison of historical 5% and 1% stationary GEV and best nonstationary GEV return levels. Gaps represent missing data.

Table 1: Fitted parameter values and skill scores for the best nonstationary GEV model.

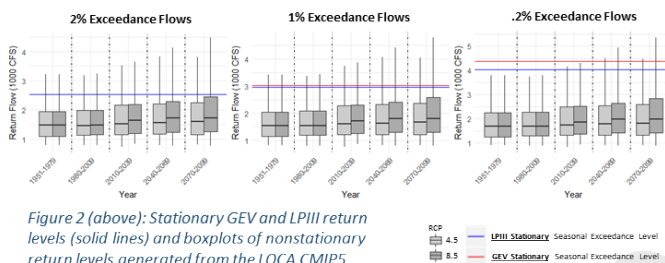


Figure 2 (above): Stationary GEV and LPIII return levels (solid lines) and boxplots of nonstationary return levels generated from the LOCA CMIP5 ensemble of covariates. Results cover two RCP scenarios.

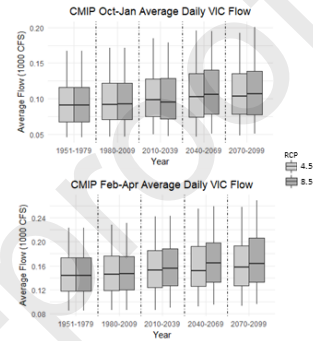


Figure 3 (above): Boxplots of LOCA CMIP5 ensemble covariates. Results cover two RCP scenarios.

Exceedance Probability	Stationary Models		Return Flow (1000 CFS)									
	GEV	LPIII	RCP	Nonstationary GEV Model - GCM Ensemble Percentile			Nonstationary GEV Model - GCM Ensemble Percentile			Nonstationary GEV Model - GCM Ensemble Percentile		
				1980-2009	2040-2069	2070-2099	1980-2009	2040-2069	2070-2099	1980-2009	2040-2069	2070-2099
20%	1.3	1.3	4.5	0.9	1.1	1.5	0.9	1.2	1.6	0.9	1.2	1.7
			8.5	0.9	1.1	1.5	1.0	1.3	1.7	1.0	1.3	1.8
10%	1.6	1.6	4.5	1.0	1.2	1.7	1.0	1.3	1.8	1.0	1.3	1.9
			8.5	1.0	1.3	1.7	1.1	1.5	1.9	1.1	1.5	2.0
4%	2.1	2.1	4.5	1.1	1.4	1.9	1.1	1.5	2.1	1.1	1.5	2.1
			8.5	1.1	1.4	1.9	1.2	1.6	2.2	1.2	1.6	2.3
1%	3.0	3.0	4.5	1.2	1.5	2.1	1.2	1.6	2.3	1.2	1.7	2.4
			8.5	1.2	1.6	2.1	1.3	1.8	2.4	1.3	1.8	2.6

Table 2 (above): Stationary GEV and LPIII return levels and best nonstationary GEV return levels. Percentiles for nonstationary return levels represent the percentile return level from the LOCA CMIP5 ensemble within the specified time period. Results cover two RCP scenarios.

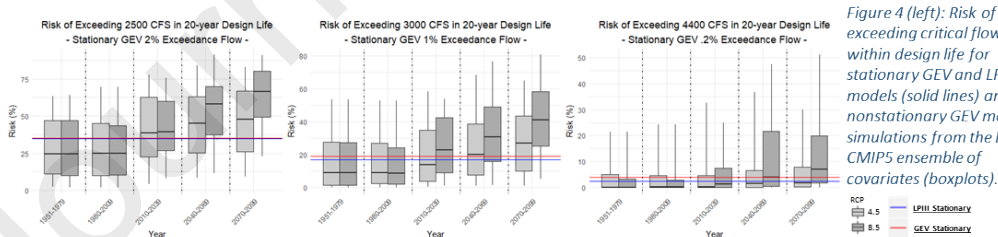


Figure 4 (left): Risk of exceeding critical flows within design life for stationary GEV and LPIII models (solid lines) and best nonstationary GEV model simulations from the LOCA CMIP5 ensemble of covariates (boxplots).

Declarations of interest: none

Journal Pre-proofs